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Radioactivity Exploration from the Arctic to Antarctica.  
Part 5: The Tundra-94 expedition

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Abstract

The joint Swedish-Russian “Tundra Ecology-94” expedition during 1994 used the large Russian ice-breaking research vessel R/V Akademik Fedorov a platform and went along a coastline of 3500 km— from the Kola Peninsula 10°E to Kolyuchinskaya Bay 173°E. Radioactivity in air, seawater and sediment was explored at various locations along the route.

The average of ⁷Be activity concentration in air over the Arctic Ocean was found to be only about 0.6 mBq.m⁻³ in air close to the Siberian coast-line, however, it was as high as 11 mBq.m⁻³. The activity concentration of ²¹⁰Pb in the air over the Arctic Ocean varies between 37 – 176 μBq.m⁻³. In the air close to the Siberian coastline 71°N 84°E, however, the activity concentration of ²¹⁰Pb in the air was much higher, about 2500 μBq.m⁻³.

Anthropogenic radioactivity in the Arctic originate from nuclear weapons fallout, release from nuclear fuel reprocessing plant, and from the Chernobyl accident. The minimum values of the ¹³⁷Cs activity concentration water along the route of the Tundra were found in South-eastern Barents Sea: 5.3 Bq.m⁻³ of surface-water, and of bottom-water 6.4 Bq.m⁻³. Maximum values were found in the Western Laptevsea: 12.8 Bq.m⁻³ of surface-water, and of bottom-water 5.1 Bq.m⁻³. East of 150 °E the ¹³⁴Cs / ¹³⁷Cs ratios are less than 0.003, indicating that less than 6% of the ¹³⁷Cs originated from the Chernobyl accident.

¹³⁷Cs levels are reduced to values of about 1.4 Bq.m⁻³ in the low salinity water near the mouths of the Ob and Yenisey Rivers. The ¹³⁴Cs / ¹³⁷Cs activity ratio of 0.014 in the freshwater indicates that the Chernobyl component in the river systems is the same (30%) as in the marine waters west of 150 °E.

In surface water the ⁹⁰Sr activity concentration range from 2 to 4 Bq.m⁻³, Maximum values about 3.5 Bq.m⁻³ were found between 100-140 °E. But east of 150 °E the values decreased to about 0.5 Bq.m⁻³ at 170 °E. In bottom water the ⁹⁰Sr activity concentration range from 1.5 at 40 °E to maximum values about 4 Bq.m⁻³ between 100-120 °E. The measured ⁹⁰Sr/¹³⁷Cs ratios in surface water close to a value of 0.14 over a wide range of stations from the Barents to the Laptev Seas. The ¹²⁹I concentration in sea-water along the route of the Tundra expedition decrease from about 20·10^{11} atoms.l⁻¹ at 40 °E, to about 1·10^{11} atoms.l⁻¹ east of 160 °E.

The ²³⁹,²⁴⁰Pu activity concentration in surface seawater decrease from about 10 mBq.m⁻³ to about 1 mBq.m⁻³ east of 160 °E. In bottom seawater it is more evenly distributed between 10⁻⁴ mBq.m⁻³, with minimum at 60-80 °E and maxima at 40°E and 160 °E. Measured ²³⁹Pu/²³⁹,²⁴⁰Pu activity ratios in the water column yield no evidence of any leakage of plutonium from dumped nuclear wastes in the Kara and Barents Seas.

Keywords: Tundra Ecology-94, Akademik Fedorov, ⁷Be, air, Arctic Ocean, Siberian coast-line, ¹³⁷Cs, ¹³⁴Cs / ¹³⁷Cs-ratio, Chernobyl accident, surface water, ⁹⁰Sr, ¹²⁹I, ²³⁹,²⁴⁰Pu, ²³⁸Pu/²³⁹,²⁴⁰Pu-activity ratio.

A. Introduction

The joint Swedish-Russian “Tundra Ecology-94” expedition during 1994 along a coastline of 3500 km—from the Kola Peninsula 10°E to Kolyuchinskaya Bay 173°E, used the large Russian ice-breaking research vessel R/V Akademik Fedorov a platform (Figure 5-1). In Table 5-1 is given locations of the various places where we were transferred a shore with helicopters (Figure 5-2).
5-6) or escorted by on Russian atomic Icebreaker (Figure 5-2). In Figure 5-3 is given a diagram of the route of the expedition.

Figure 5-1a.
R/V Akademik Fedorov in the harbour of Gothenburg
loading of the Tundra-94 expedition

Figure 5-1b.
R/V Akademik Fedorov, ready to leave for the
Arctic

Table 5-1
Expedition route and research sites of the "Tundra Ecology-94" expedition

<table>
<thead>
<tr>
<th>Visiting date 1</th>
<th>On return date 2</th>
<th>Site No.</th>
<th>Name</th>
<th>Position</th>
</tr>
</thead>
<tbody>
<tr>
<td>06-04</td>
<td>09-08</td>
<td></td>
<td>Gothenburg</td>
<td>57.43°N 11.98°E</td>
</tr>
<tr>
<td>06-08&gt;&gt;09</td>
<td>09-02</td>
<td></td>
<td>Murmansk</td>
<td>68.57°N 44.10°E</td>
</tr>
<tr>
<td>06-10</td>
<td>08-31</td>
<td>1(27)</td>
<td>Kachkovsky Bay. Kola Peninsula</td>
<td>67.30°N 41.00°E</td>
</tr>
<tr>
<td>06-12&gt;&gt;13</td>
<td>08-29&gt;&gt;30</td>
<td>2(26)</td>
<td>North-Eastern Kanin Peninsula</td>
<td>68.15°N 6.00°E</td>
</tr>
<tr>
<td>06-14&gt;&gt;15</td>
<td>08-26&gt;&gt;28</td>
<td>3(25)</td>
<td>Kolguyev Island</td>
<td>69.15°N 50.00°E</td>
</tr>
<tr>
<td>06-15&gt;&gt;16</td>
<td>08-25&gt;&gt;26</td>
<td>4(24)</td>
<td>Pechora Bay</td>
<td>68.5°N 54.00°E</td>
</tr>
<tr>
<td>06-17&gt;&gt;20</td>
<td>08-22&gt;&gt;08-23</td>
<td>5(23)</td>
<td>Western Yamal Peninsula</td>
<td>70.45°N 67.00°E</td>
</tr>
<tr>
<td>06-21&gt;&gt;22</td>
<td>08-20&gt;&gt;21</td>
<td>6 (22)</td>
<td>Belyi Island. Northern Yamal Peninsula</td>
<td>73.00°N 70.00°E</td>
</tr>
<tr>
<td>06-22&gt;&gt;23</td>
<td>08-18&gt;&gt;19</td>
<td></td>
<td>Dickson</td>
<td></td>
</tr>
<tr>
<td>06-23</td>
<td></td>
<td></td>
<td>Arctic Institute Islands</td>
<td>75.00°N, 82.00°E</td>
</tr>
<tr>
<td>06-24&gt;&gt;26</td>
<td>08-15&gt;&gt;17</td>
<td>8(21)</td>
<td>North West of Taymyr Peninsula</td>
<td>76.00°N 94.00°E</td>
</tr>
<tr>
<td>06-27&gt;&gt;28</td>
<td>08-13&gt;&gt;14</td>
<td>9(20)</td>
<td>Chelyuskin Peninsula</td>
<td>77.20°N 102.00°E</td>
</tr>
<tr>
<td>08-29&gt;&gt;30</td>
<td>08-10&gt;&gt;11</td>
<td>10(19)</td>
<td>North-east of Taymyr Peninsula</td>
<td>76.00°N 112.00°E</td>
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<tr>
<td>07-03&gt;&gt;05</td>
<td></td>
<td></td>
<td>Khatanga rotation point 1</td>
<td>74.00°N 110.00°E</td>
</tr>
<tr>
<td>07-05&gt;&gt;06</td>
<td></td>
<td>11</td>
<td>Olenekskiy Bay</td>
<td>73.15°N 120.00°E</td>
</tr>
<tr>
<td>08-08&gt;&gt;08-08</td>
<td></td>
<td></td>
<td>Tiksi rotation point 2</td>
<td>74.00°N 110.00°E</td>
</tr>
<tr>
<td>08-03&gt;&gt;08-04</td>
<td></td>
<td>12</td>
<td>Yana Delta</td>
<td>71.30°N 136.00°E</td>
</tr>
<tr>
<td>07-10&gt;&gt;11</td>
<td>07-31&gt;&gt;08-02</td>
<td>13</td>
<td>New Siberian Islands</td>
<td>75.00°N 140.00°E</td>
</tr>
<tr>
<td>07-14&gt;&gt;15</td>
<td></td>
<td>14</td>
<td>Lopatka Peninsular, N-W Indigirka</td>
<td>71.45°N 149.00°E</td>
</tr>
<tr>
<td>07-17&gt;&gt;18</td>
<td></td>
<td>15</td>
<td>North-east of Kolyma Delta</td>
<td>71.45°N 158.00°E</td>
</tr>
<tr>
<td>07-20&gt;&gt;21</td>
<td></td>
<td>16</td>
<td>Ayon Island</td>
<td>69.50°N 168.00°E</td>
</tr>
<tr>
<td>07-22&gt;&gt;26</td>
<td></td>
<td>17</td>
<td>South-western Wrangels Island</td>
<td>70.50°N 179.00°E</td>
</tr>
<tr>
<td>07-25</td>
<td>point of return</td>
<td>18</td>
<td>Kolyuchinskaya Bay</td>
<td>67.00°N 173.45°E</td>
</tr>
</tbody>
</table>
Figure 5-2.
The Russian atomic icebreaker approaching To assist Akademik Fedorov in the heavy ice.

![Image of the Russian atomic icebreaker approaching To assist Akademik Fedorov in the heavy ice.](image)

Figure 5-3.
Route of the Tundra-94 expedition with R/V Akademik Fedorov along the Siberian Coastline, with a helicopter tour to Wrangles Island.

B. Material and Methods.

B1. Air sampling

Air filter samples were taken by using an Andersen sampler with the capacity of 100 m$^3$h$^{-1}$ (filter size 0.25x0.25 m membrane filter). An air volume of about 1 500 m$^3$ were collected at a
rate of 100 m$^3$h$^{-1}$ during each sampling occasion. The Andersen sampler was placed close together with a FOA transportable reference high volume air sampler (filter size 0.56x0.56 m. microsorban filter. 1 100 m$^3$h$^{-1}$), previously taken part in an inter calibration of samplers (Vintersved, 1994). The filters were stored on board and then taken the institute at Lund for analysis. The results of the Anderson sampler were normalized to those of the calibrated FOA sampler based on $^{7}$Be measurements. FOA nowadays FOI, is a Swedish research institute in the areas of defence and security.

![Figure 5-4. The FOA air sampler](image)

**B2. Analysis of the air filters**

The filters were measured for $^{7}$Be by gamma spectrometry using a high performance Germanium detector (HpGe Canberra). After adding $^{209}$Po as radiochemical yield determinant, the samples were wet-ashed by using a mixture of concentrated nitric and per-chloric acids. Polonium was spontaneously deposited on nickel discs, and measured by alpha spectrometry, using surface ion implanted silicon detectors.

Remaining traces of polonium, were removed by anion exchange. The solution was then stored for about 8 months to allow equilibrium in-growth of $^{210}$Po from $^{210}$Pb. The activity of ingrown $^{210}$Po, was then analysed as described above, and finally the activity concentrations of $^{210}$Po and $^{210}$Pb in air were calculated.

**B3. Water sampling**

Large volume (200 litre) water samples, were taken from the water cooling system of the ship, and collected in special vessels for precipitation of $^{137}$Cs and $^{239+240}$Pu. Continuous sampling of caesium also took place with a separate pump and a pipe hanging from the rail of the ship to about 2 m depth. An in-line system with filters impregnated with Copper-Ferro-Cyanide ($\text{Cu}_2\text{Fe(CN)}_6$) was used to collect Cesium isotopes from the seawater. The filters were dried and brought to Lund for radiochemical analysis. After ashing the filters at 420 °C, the residues were analysed for $^{134}$Cs and $^{137}$Cs by using a Ge (Li) gamma spectrometer.
Figure 5-5.
Interior of the water laboratory container, with the two 200 litre precipitation vessels to the right and the cartridge filters on the wall to the left.

Figure 5-6.
Bertil Persson and Kjell-Åke Carlsson landed on the tundra after a tour with the Russian helicopter in the back.
Figure 5-6a. View of the tundra

Figure 5-6b. Sampling of the tundra

Figure 5-6a. Closer view

Figure 5-7a. Chelyuskin Peninsula 77.20°N; 102.00°E
Summer lake of the waste accumulated during the winter.

Figure 5-7a. Road at Chelyuskin Peninsula

Figure 5-7a. A mound of flat stones raised by Adolf Erik Nordenskiöld’s expedition in 1878 as a memorial of the visit.

Figure 5-7b. An anchor left by Adolf Erik Nordenskiöld’s expedition in 1878.
C. Results

C1. $^{210}$Pb and $^7$Be in air 1994- June 08 > September 08

In the Figures 5-8a and b are given the activity concentrations of $^{210}$Pb ($\mu$Bq.m$^{-3}$), and $^7$Be (mBq.m$^{-3}$) in air, measured during 1994-June 8 < September 8 at the joint Swedish-Russian Tundra Ecology-94 expedition.

**Figure 5-8 a**
Longitudinal distribution of $^{210}$Pb air concentration ($\mu$Bq.m$^{-3}$) during 1994-June 8 < September 8 at the Tundra Ecology-94 expedition. Predicted values in red.

**Figure 5-8 b** Longitudinal distribution of $^7$Be air concentration (mBq.m$^{-3}$) during 1994-June 8 < September 8 and the route of the Tundra Ecology-94 expedition. Predicted values in red.

Equations of the PLS models for the air concentrations displayed in red in Figure 5-8:

$$C_{^7Be} = 24.65 - 0.207 \cdot \text{Latitude} - 0.016 \cdot \text{Longitude}; \ [\text{mBq.m}^{-3}]$$
Goodness of fit statistics (Variable $C_{^7Be}$): $R^2 = 0.036$

$$C_{^{210}Pb} = 6992 - 44.45 \cdot \text{Latitude} - 14.61 \cdot \text{Longitude}; \ [\mu\text{Bq.m}^{-3}]$$
Goodness of fit statistics (Variable $C_{^{210}Pb}$): $R^2 = 0.176$
Table 5-2a
Average air concentrations of $^{7}$Be and $^{210}$Pb measured during 1994-June 8 > September 8 at the joint Swedish-Russian Tundra Ecology-94 expedition.

<table>
<thead>
<tr>
<th>Isotope</th>
<th>Date</th>
<th>Average</th>
<th>SD</th>
<th>SE</th>
<th>Site</th>
</tr>
</thead>
<tbody>
<tr>
<td>$^{7}$Be</td>
<td>1994-0605 &gt; 0719</td>
<td>11.4</td>
<td>9.0</td>
<td>3.2</td>
<td>N</td>
</tr>
<tr>
<td>$^{210}$Pb</td>
<td>1994-0605 &gt; 0719</td>
<td>2373</td>
<td>1029</td>
<td>364</td>
<td>μBq/m³</td>
</tr>
<tr>
<td>$^{7}$Be</td>
<td>1994-07-19 &gt; 0908</td>
<td>7.2</td>
<td>5.4</td>
<td>2.0</td>
<td>N</td>
</tr>
<tr>
<td>$^{210}$Pb</td>
<td>1994-07-19 &gt; 0908</td>
<td>2712</td>
<td>2854</td>
<td>1079</td>
<td>μBq/m³</td>
</tr>
</tbody>
</table>

During the Swedish-Russian Tundra Ecology-94 expedition along the Siberian coastline, the average air concentrations of $^{7}$Be and $^{210}$Pb measured during May-July were 11±3 and 2.4±0.4 mBq.m⁻³ respectively and during July-September they were 7.2±2 and 2.7±1.1 mBq.m⁻³ respectively.

Table 5-2 b
Ratios of average air concentrations of $^{7}$Be and $^{210}$Pb measured during 1994-June 8 > September 8 at the joint Swedish-Russian Tundra Ecology-94 expedition.

<table>
<thead>
<tr>
<th>Isotope ratio</th>
<th>Date</th>
<th>Average</th>
<th>SE</th>
</tr>
</thead>
<tbody>
<tr>
<td>$^{7}$Be / $^{210}$Pb</td>
<td>1994-0605 &gt; 0719</td>
<td>5</td>
<td>2</td>
</tr>
<tr>
<td>$^{7}$Be / $^{210}$Pb</td>
<td>1994-07-19 &gt; 0908</td>
<td>3</td>
<td>1</td>
</tr>
</tbody>
</table>

C2. $^{137}$Cs activity concentration in seawater
The $^{137}$Cs activity concentration water along the route of the Tundra expedition is shown in figures 5-9a and 9b respectively. The minimum values were found in South-Eastern Barents Sea: 5.3 Bq.m⁻³ of surface-water, and of bottom-water 6.4 Bq.m⁻³. Maximum values were found in the Western Laptevsea: 12.8 Bq.m⁻³ of surface-water, and of bottom-water 5.1 Bq.m⁻³.

![Figure 5-9a](image_url)

**Figure 5-9a.** Longitudinal distribution of $^{137}$Cs activity concentration in surface seawater along the route of Tundra expedition. Predicted values in read...
Equation of the PLS model for activity concentration in surface water (SW), displayed in red in Figure 5-9:

$$^{\text{SW}}C_{137\text{Cs}} = -44.93 + 0.756\cdot(\text{Latitude}^\circ\text{N}) - 0.035\cdot(\text{Longitude}^\circ\text{E}); \quad \text{[Bq.m}^{-3}\text{]}$$

Goodness of fit statistics: $R^2 = 0.312$

Equation of the PLS model for bottom water (BW):

$$^{\text{BW}}C_{137\text{Cs}} = -2.005 + 0.071\cdot(\text{Latitude}^\circ\text{N}) + 4.49\cdot10^{-5}\cdot(\text{Longitude}^\circ\text{E}); \quad \text{[Bq.m}^{-3}\text{]}$$

Goodness of fit statistics (Variable $^{\text{BW}}C_{137\text{Cs}}$ bottom water): $R^2 = 0.076$

Figure 5-10a. View of the Arctic sea

Figure 5-10b. View of the Arctic sea

Figure 5-10c. View of the Arctic sea

C3. $^{134}\text{Cs}/^{137}\text{Cs}$ activity ratio in sea water

Figure 5-11a.
Longitudinal distribution of $^{134}\text{Cs}/^{137}\text{Cs}$ activity ratio in surface sea-water along the route of Tundra expedition. Predicted values in red
Figure 5-11b.
Longitudinal distribution of $^{134}$Cs/$^{137}$Cs activity ratio in bottom sea water along the route of Tundra expedition. Predicted values in read

Equation of the PLS model for $^{134}$Cs/$^{137}$Cs ratio in surface water (SW):

$\left(\frac{^{134}\text{Cs}}{^{137}\text{Cs}}\right)_{\text{SW}} = -0.011 + 3.98\times10^{-4}\cdot(\text{Latitude } ^\circ\text{N}) - 4.85\times10^{-5}\cdot(\text{Longitude } ^\circ\text{E})$

Goodness of fit statistics (Variable $^{134}$Cs/$^{137}$Cs Surface water): $R^2 = 0.447$

Equation of the model for bottom water (BW):

$\left(\frac{^{134}\text{Cs}}{^{137}\text{Cs}}\right)_{\text{BW}} = -0.042 + 8.89\times10^{-4}\cdot(\text{Latitude } ^\circ\text{N}) – 9.37\times10^{-5}\cdot(\text{Longitude } ^\circ\text{E})$

Goodness of fit statistics (Variable $^{134}$Cs/$^{137}$Cs Bottom water): $R^2 = 0.913$

C4. $^{90}$Sr activity concentration in seawater

The $^{90}$Sr activity concentration in seawater along the route of the Tundra expedition is shown in figures 5-13a and b respectively. In surface water the $^{90}$Sr activity concentration range from 2 to 4 Bq.m$^{-3}$. Maximum values about 3.5 Bq.m$^{-3}$ were found between 100-140 $^\circ$E. But east of 150 $^\circ$E, the values decreased to about 0.5 Bq.m$^{-3}$ at 170 $^\circ$E.

Figure 5-12a Longitudinal distribution of the $^{90}$Sr activity concentration in surface seawater along the route of Tundra expedition. Predicted values in read
In bottom water the $^{90}$Sr activity concentration range from 1.5 at 40 °E to maximum values about 4 Bq.m$^{-3}$ between 100-120 °E.

Equation of the PLS model for $^{90}$Sr activity concentration in Surface water (SW):

$^{SW}C_{90Sr} = -7.910 + 0.153 \cdot (\text{Latitude}^\circ\text{N}) + 9.6 \cdot 10^{-5} \cdot (\text{Longitude}^\circ\text{E});$ [Bq.m$^{-3}$]

Goodness of fit statistics: $R^2 = 0.16$

Equation of the PLS model $^{90}$Sr activity concentration in Bottom Water (BW):

$^{BW}C_{90Sr} = -0.873 + 0.052 \cdot (\text{Latitude}^\circ\text{N}) + 0.0025 \cdot (\text{Longitude}^\circ\text{E});$ [Bq.m$^{-3}$]

Goodness of fit statistics: $R^2 = 0.21$

---

**Figure 5-12b** Longitudinal distribution of the $^{90}$Sr activity concentration in surface and bottom seawater along the route of Tundra expedition. Predicted values in read

**Figure 5-13a** Longitudinal distribution of the $^{90}$Sr/$^{137}$Cs activity ratio in surface seawater along the route of Tundra expedition.
Figure 5b Longitudinal distribution of the $^{90}\text{Sr} / ^{137}\text{Cs}$ activity ratio in bottom seawater and the route of Tundra expedition.

C5. $^{129}\text{I}$ concentration in seawater

The $^{129}\text{I}$ concentration in sea-water along the route of the Tundra expedition is shown in Figure 5-14. The concentration decrease from about $20 \cdot 10^{11}$ atoms.$\text{l}^{-1}$ ($2 \cdot 10^{15}$ atoms.$\text{m}^{-3}$), $\approx$ 3 picomolar, to about $1 \cdot 10^{11}$ atoms.$\text{l}^{-1}$ east of 160 $^\circ$E.

Figure 5-14 Longitudinal distribution of $^{129}\text{I}$ concentration in surface seawater and the route of Tundra expedition. Predicted values in read.
PLS Model parameters for $^{129}$I concentration in surface seawater:

$$SWC_{^{129}I} = -25.529 + 0.580 \cdot (\text{Latitude} \, ^{\circ}N) - 0.065 \cdot (\text{Longitude} \, ^{\circ}E); \left[10^{14} \text{ atoms.m}^{-3}\right]$$

Goodness of fit statistics: $R^2 = 0.311$

C6. $^{239+240}$Pu activity concentration in seawater

The $^{239+240}$Pu activity concentration in surface and bottom seawater along the route of Tundra expedition are given in Figure 3-15a and b respectively. In surface seawater the $^{239+240}$Pu activity concentration decrease from about 10 mBq.m$^{-3}$ to about 1 mBq.m$^{-3}$ east of 160 $^{\circ}$E.

![Figure 5-15a. $^{239+240}$Pu activity concentration in surface seawater along the route of Tundra expedition. Predicted values in read.](image)

In bottom seawater the $^{239+240}$Pu activity concentration is more evenly distributed between 10 - 4 mBq.m$^{-3}$, with the minimum at 60-80 $^{\circ}$E and maxima at 40$^{\circ}$E and 160 $^{\circ}$E.

![Figure 5-15b. $^{239+240}$Pu activity concentration in bottom seawater and the route of Tundra expedition.](image)
Equation of the model PLS model for activity concentration of $^{239,240}$Pu in Surface water (SW):

$$SW_{C_{239,240Pu}} = 5.948 + 0.028 \cdot (\text{Latitude}^\circ \text{N}) - 0.043 \cdot (\text{Longitude}^\circ \text{E}) \; [\text{mBq.m}^{-3}]$$

Goodness of fit statistics (Variable $SW_{C_{239,240Pu}}$): $R^2 = 0.471$

Equation of the PLS model for $^{239,240}$Pu Bottom water (BW):

$$BW_{C_{239,240Pu}} = 30.96 - 0.331 \cdot (\text{Latitude}^\circ \text{N}) - 0.0028 \cdot (\text{Longitude}^\circ \text{E}) \; [\text{mBq.m}^{-3}]$$

Goodness of fit statistics (Variable $BW_{C_{239,240Pu}}$): $R^2 = 0.432$

C7. $^{137}$Cs and $^{239+240}$Pu activity concentration in sediment

The integrated sediment activity of $^{234+240}$Pu was measured in samples taken at sampling sites specified in Table 5-3, and the results are displayed in Figure 5-16a and b.

Table 5-3

<table>
<thead>
<tr>
<th>Station</th>
<th>Latitude</th>
<th>Longitude</th>
<th>Water depth</th>
<th>Activity 137Cs</th>
<th>$^{234+240}$Pu</th>
</tr>
</thead>
<tbody>
<tr>
<td></td>
<td>$^\circ$N</td>
<td>$^\circ$E</td>
<td>m</td>
<td>[Bq/m²]</td>
<td>[Bq/m²]</td>
</tr>
<tr>
<td>9</td>
<td>70</td>
<td>14</td>
<td>66</td>
<td>17</td>
<td>20</td>
</tr>
<tr>
<td>14</td>
<td>76</td>
<td>11</td>
<td>93</td>
<td>34</td>
<td>55</td>
</tr>
<tr>
<td>18</td>
<td>75</td>
<td>8</td>
<td>129</td>
<td>50</td>
<td>58</td>
</tr>
<tr>
<td>21</td>
<td>74</td>
<td>50</td>
<td>137</td>
<td>32</td>
<td>20</td>
</tr>
<tr>
<td>28</td>
<td>70</td>
<td>16</td>
<td>170</td>
<td>26</td>
<td>30</td>
</tr>
</tbody>
</table>

Figure 5-16a

Measured and predicted, integrated sediment activity of $^{234+240}$Pu at specific sampling stations (see below). Predicted values in read.
Equation of the model for predicted integrated sediment activity, $S_{\text{ED}}^A$, of $^{239+240}\text{Pu}$ [Bq/m²] in sediment:

$$S_{\text{ED}}^A_{239+240\text{Pu}} = -115.2 - 0.215 \cdot (\text{Longitude}^\circ\text{E}) + 2.23 \cdot (\text{Latitude}^\circ\text{N})$$

Goodness of fit statistics ($S_{\text{ED}}^A_{239+240\text{Pu}}$): $R^2 = 0.852$

Equation of the PLS model for predicted integrated sediment activity, $S_{\text{ED}}^A$, of $^{137}\text{Cs}$ [Bq/m²] in sediment ($S_{\text{ED}}$):

$$S_{\text{ED}}^A_{137\text{Cs}} = -473.3 - 1.486 \cdot (\text{Longitude}^\circ\text{E}) + 14.66 \cdot (\text{Latitude}^\circ\text{N}); [\text{Bq.m}^{-2}]$$

Goodness of fit statistics (Variable $137\text{Cs}$ [Bq/m²]): $R^2 = 0.069$

D. Discussions

D1. Beryllium-7 activity concentrations in the Arctic air

The activity-concentration of $^7\text{Be}$ in air in the Arctic air as summarized in Table 5-4 varies between 2 - 4.9 mBq.m³ with average $2.8\pm0.3$ mBq.m³ (Buraglio et al., 2001, Kulan, 2006, Paatero and Hatakka, 2000, Baskaran and Shaw, 2001, Dibb and Jaffrezo, 1993). The average of $^7\text{Be}$ activity concentration in air over the Arctic Ocean was, however, only about 0.6 mBq.m³. In contrast the activity concentration of $^7\text{Be}$ in air close to the Siberian coast-line as high as 11 mBq.m³ (Persson, 2013).
Table 5-4. Summary of atmospheric $^7$Be concentrations in Arctic and sub-Arctic air

<table>
<thead>
<tr>
<th>Time</th>
<th>Location</th>
<th>Lat</th>
<th>Long</th>
<th>Be-7 Arithm. Mean mBq.m$^{-3}$</th>
<th>SD</th>
</tr>
</thead>
<tbody>
<tr>
<td>910728-0906</td>
<td>Arctic Ocean</td>
<td>82.07</td>
<td>51.00</td>
<td>0.62</td>
<td>0.52</td>
</tr>
<tr>
<td>910907-1004</td>
<td>Arctic Ocean</td>
<td>84.36</td>
<td>-2.32</td>
<td>0.51</td>
<td>0.33</td>
</tr>
<tr>
<td>940605-0908</td>
<td>N Siberian coast</td>
<td>71</td>
<td>84</td>
<td>11.4</td>
<td>9.0</td>
</tr>
</tbody>
</table>

<table>
<thead>
<tr>
<th>Time</th>
<th>Location</th>
<th>Lat</th>
<th>Long</th>
<th>Pb-210 Average µBq.m$^{-3}$</th>
<th>SD</th>
<th>Po-210 Average µBq.m$^{-3}$</th>
<th>SD</th>
<th>Reference</th>
</tr>
</thead>
<tbody>
<tr>
<td>2000</td>
<td>Uppsala. Sweden</td>
<td>59.88</td>
<td>17.63</td>
<td>4.7</td>
<td>2.3</td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>1972-1995</td>
<td>Sweden</td>
<td>59.88</td>
<td>17.63</td>
<td>4.8</td>
<td>2.4</td>
<td></td>
<td></td>
<td>(Kulan, 2006)</td>
</tr>
<tr>
<td>1972-2003</td>
<td>Sweden, Kiruna</td>
<td>67.84</td>
<td>20.34</td>
<td>1.9</td>
<td>1.0</td>
<td></td>
<td></td>
<td>(Kulan, 2006)</td>
</tr>
<tr>
<td>1972-2003</td>
<td>Sweden, Grindsjön</td>
<td>59.07</td>
<td>17.82</td>
<td>2.3</td>
<td>1.2</td>
<td></td>
<td></td>
<td>(Kulan, 2006, Aldahan et al., 2008)</td>
</tr>
<tr>
<td>1972-2003</td>
<td>Sweden, Ljungbyhed</td>
<td>56.08</td>
<td>13.23</td>
<td>2.5</td>
<td>1.3</td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>1995-1997</td>
<td>Finland Sodankyla</td>
<td>67.37</td>
<td>26.65</td>
<td>2.5</td>
<td>2.0</td>
<td></td>
<td></td>
<td>(Paatero, 2000)</td>
</tr>
<tr>
<td>1996</td>
<td>Alaska USA Poker Flat</td>
<td>65.13</td>
<td>-147.48</td>
<td>3.0</td>
<td>2.0</td>
<td></td>
<td></td>
<td>(Baskaran and Shaw, 2001)</td>
</tr>
<tr>
<td>1996</td>
<td>Alaska USA Eagle</td>
<td>65.9</td>
<td>-141.20</td>
<td>2.2</td>
<td>1.0</td>
<td></td>
<td></td>
<td>(Baskaran and Shaw, 2001)</td>
</tr>
<tr>
<td>1988-1989</td>
<td>Dye3</td>
<td>65.18</td>
<td>43.82</td>
<td>2.6</td>
<td>1.1</td>
<td></td>
<td></td>
<td>(Dibb and Jaffrezo, 1993)</td>
</tr>
<tr>
<td>1988-1990</td>
<td>Barrow</td>
<td>71.30</td>
<td>-156.77</td>
<td>1.9</td>
<td>1.1</td>
<td></td>
<td></td>
<td>(Dibb and Jaffrezo, 1993)</td>
</tr>
<tr>
<td>1988-1991</td>
<td>Kap Toban</td>
<td>70.42</td>
<td>-21.97</td>
<td>2.4</td>
<td>1.3</td>
<td></td>
<td></td>
<td>(Dibb and Jaffrezo, 1993)</td>
</tr>
<tr>
<td>1988-1992</td>
<td>Nord</td>
<td>81.36</td>
<td>-16.40</td>
<td>2.5</td>
<td>1.4</td>
<td></td>
<td></td>
<td>(Dibb and Jaffrezo, 1993)</td>
</tr>
<tr>
<td>1988-1993</td>
<td>Thule</td>
<td>77.50</td>
<td>-69.33</td>
<td>3.7</td>
<td>1.9</td>
<td></td>
<td></td>
<td>(Dibb and Jaffrezo, 1993)</td>
</tr>
<tr>
<td>1997-2004</td>
<td>Summit, Greenland</td>
<td>72.575</td>
<td>-27.55</td>
<td>2.0</td>
<td>0.5</td>
<td></td>
<td></td>
<td>(Dibb, 2007)</td>
</tr>
</tbody>
</table>

$^a$(Persson, 2013).

D2. $^{210}$Pb activity concentrations in the Arctic air

Observations of the activity concentration of $^{210}$Pb in the air over the Arctic ocean as summarized in Table 5-5, varies between 37 – 176 µBq.m$^{-3}$ (Persson and Holm, 2013, McNeary and Baskaran, 2003, Dibb and Jaffrezo, 1993, Dibb, 2007, Paatero et al., 2003, Samuelsson et al., 1986). In 1991 we found the average activity concentration of $^{210}$Pb over the Arctic Ocean to be 40±4 µBq.m$^{-3}$. In the air close to land masses the activity concentration of $^{210}$Pb in the air increase to 269- 2712 µBq.m$^{-3}$ (McNeary and Baskaran, 2003, Baskaran and Shaw, 2001, Dibb and Jaffrezo, 1993); with the highest values of about 2500 µBq.m$^{-3}$ at the Siberian coastline (Persson and Holm, 2013).

Table 5-5. Activity concentrations (µBq.m$^{-3}$) of $^{210}$Pb recorded at different locations during the Tundra-94 expedition.

<table>
<thead>
<tr>
<th>Time</th>
<th>Location</th>
<th>Lat °N</th>
<th>Long °E</th>
<th>Pb-210 Average µBq.m$^{-3}$</th>
<th>SD</th>
<th>Po-210 Average µBq.m$^{-3}$</th>
<th>SD</th>
<th>Reference</th>
</tr>
</thead>
<tbody>
<tr>
<td>940605&gt;0703</td>
<td>Siberian Tundra</td>
<td>71</td>
<td>84</td>
<td>2373</td>
<td>364</td>
<td>2044</td>
<td>870</td>
<td>This work</td>
</tr>
<tr>
<td>940704&gt;0908</td>
<td>Siberian Tundra</td>
<td>71</td>
<td>84</td>
<td>2712</td>
<td>1079</td>
<td>2336</td>
<td>994</td>
<td>This work</td>
</tr>
</tbody>
</table>
D3. $^{137}$Cs activity distribution

The minimum values of the $^{137}$Cs activity concentration water along the route of the Tundra were found in South-eastern Barents Sea: $5.3 \text{ Bq.m}^{-3}$ of surface-water, and of bottom-water $6.4 \text{ Bq.m}^{-3}$. Maximum values were found in the Western Laptev sea: $12.8 \text{ Bq.m}^{-3}$ of surface-water, and of bottom-water $5.1 \text{ Bq.m}^{-3}$.

East of 150°E the $^{134}$Cs / $^{137}$Cs ratios are less than 0.003, indicating that less than 6% of the $^{137}$Cs originated from the Chernobyl accident.

The $^{134}$Cs / $^{137}$Cs activity ratio of 0.014 in the freshwater indicates that the Chernobyl component in the river systems is the same (30%) as in the marine waters.

D4. $^{90}$Sr activity distribution

The relative magnitudes of $^{90}$Sr inputs to the Arctic Ocean differ from those for $^{137}$Cs for the same sources. The $^{137}$Cs/$^{90}$Sr activity ratio of 35 reported for Chernobyl fallout was sufficiently high that $^{90}$Sr inputs from this source can be considered to be negligible (Aarkrog, 1988). The present fallout concentration in the oceans is assumed to be about 1.6 Bq/m$^3$ (Dahlgaard, 1995). Sellafield represents a major $^{90}$Sr source term, which similar to $^{137}$Cs, attained a maxima in the late 1970's and has decreased substantially since that time. An important additional source of $^{90}$Sr to the Siberian seas is associated with river runoff from fallout, discharges from nuclear reprocessing plants and inputs from accidental releases of $^{90}$Sr, such as the Khystym accident on 29 September 1957 at Mayak, USSR (Lollino et al., 2014). The greater mobility of $^{90}$Sr compared to $^{137}$Cs in freshwater environments results in reduced $^{90}$Sr residence times in soils and more rapid transport through the drainage basin to marginal seas. By using the record of reported $^{90}$Sr discharges, transport times of less than 10 y and transfer factors of 10 Bq.m$^{-3}$ per PBq.a$^{-1}$ the Sellafield contribution to Barents Sea water is estimated to be approximately 0.5 Bq.m$^{-3}$ in 1994 (Gray et al., 1995). The addition of a fallout component of approximately 1.6 Bq.m$^{-3}$ is not sufficient to give the values ($> 3 \text{ Bq.m}^{-3}$) measured in the Kara and Laptev Seas. These results suggest an additional contribution of the order of 1-2 Bq.m$^{-3}$ to $^{90}$Sr concentrations on the Siberian shelves. Contributions from riverine sources will generally only play a minor role since most the salinities are too high.

The $^{90}$Sr/$^{137}$Cs fallout ratio in seawater is approximately 0.7 while Sellafield discharge results give an average cumulative decay corrected $^{90}$Sr/$^{137}$Cs ratio of 7 between 1980 and 1992 (Gray et al., 1995, Dahlgaard et al., 1995). The measured $^{90}$Sr/$^{137}$Cs (non-Chernobyl) ratios in surface water (Figure 5-13) are also close to a value of 0.14 over a wide range of stations from the Barents to the Laptev Seas, despite the observation above that much of this signal is from Sellafield. Clearly, the $^{90}$Sr input from the Russian river systems has been sufficiently large to reduce the $^{137}$Cs/$^{90}$Sr activity ratio to values similar to fallout levels. Bottom waters show slightly higher ratios indicating a Sellafield contribution. Calculations as above applying known transfer factors and transport times reveal, however, that direct transport will not notably effect the fallout ratio. Instead the Sellafield activity must be of an older date, reflecting a longer half-life on the shelf than expected, or recirculated from the central Arctic Ocean.
D5. \(^{129}\)I distribution

The decreasing gradient in \(^{129}\)I activity, east of the Barents Sea to the Laptev Sea, reflects the general increase in \(^{129}\)I discharges since the 1980’s. The sharp decrease in \(^{129}\)I concentrations at 150 °E indicates that the front between Atlantic and Pacific origin water has been encountered. In the 1980s, this front was located over the Lomonosov Ridge, but has shifted to its present position over the Mendelyev Ridge (McLaughlin et al., 1996, Smith et al., 1998). The relatively low radionuclide levels measured in the East Siberian Sea are typical of fallout values associated with Pacific-origin water transported into this region from the Bering Sea.

D6. \(^{239}+^{240}\)Pu activity distribution

Plutonium activity concentrations and isotopic ratios, measured along the Siberian Shelf and in the Central Arctic Ocean, indicate that it mainly originates from global fallout of atmospheric nuclear weapons tests. This demonstrate that plutonium fallout of atmospheric nuclear weapons tests, deposited at mid-latitudes in the North Atlantic in the late-1950s and early-1960s, have found their way to the Arctic interior and beyond (Kershaw and Baxter, 1995, Josefsson, 1998); Herrmann, 1998 #473).

Measured \(^{238}\)Pu/\(^{239,240}\)Pu activity ratios in the water column yield no evidence of any leakage of plutonium from dumped nuclear wastes in the Kara and Barents Seas. Were leakage of plutonium to occur in the future from dumped nuclear wastes in the Kara and Barents Seas, it is likely that some of it will be transported through the Eurasian Shelf and into the Central Arctic with the Transpolar Drift, on a timescale of one to two decades, eventually exiting the Arctic through Fram Strait.

The geographical distribution of plutonium indicate that a broad peak that appears to have passed through the North Pole recently. We attribute this peak to the plutonium ‘signal’ that entered the Arctic following the period of maximal fallout deposition referred to above. The distribution is consistent with the well-established pattern of water-mass circulation in the Arctic, bearing in mind the limited number of plutonium observations available.

E. References


KULAN, A. 2006. Seasonal Be-7 and Cs-137 activities in surface air before and after the Chernobyl event. *Journal of Environmental Radioactivity*, 90, 140-150.


PERSSON, B. R. R. & HOLM, E. 2013. $^7$Be. $^{210}$Pb. and $^{210}$Po in the surface air from the Arctic to Antarctica. *INCOPoPb-2013 conference, Mangalore University*.

