

1 **Methane producing and reducing microorganisms display a high resilience to**
2 **drought in a Swedish hemi-boreal mire**

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4 **Parmentier**^{1,6}

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16

17 **Key Points:**

- 18 • Taxonomic and functional gene composition significantly changed during the drought
19 • Methane fluxes significantly reduced during drought but not in all ecotypes
20 • Specialists genera respond to drought stronger than others
21

22 Abstract

23 An increased frequency of droughts due to anthropogenic climate change can lead to considerable
24 stress for soil microorganisms and their functioning within northern peatlands. A better
25 understanding of the diversity and abundance of methane producing and reducing taxa, and their
26 functional genes, can help predict the functional potential of peatlands and how the
27 microorganisms respond to disturbances such as drought. In order to address knowledge gaps in
28 the understanding of how functional genetic diversity shifts under drought conditions, we
29 investigated a hemi boreal mire in Southern Sweden. Environmental parameters, including soil and
30 air temperature, precipitation and water table depth, as well as methane flux data were collected
31 during the summer of 2017 under typical growing conditions, and in 2018 during a drought. In
32 addition, the diversity and composition of genes encoding for methane metabolism were
33 determined using the captured metagenomics technique. During drought we observed a substantial
34 increase in air and soil temperature, reduced precipitation, and a lower water table depth.
35 Taxonomic and functional gene composition significantly changed during the drought, while
36 diversity indices, such as alpha and beta diversity, remained similar. These results indicate that
37 methane producing and reducing microbial communities, and their functional genes, displayed a
38 resilience to drought with specific genera having the ability to outcompete others under stress.
39 Furthermore, our results show that although methane emissions are substantially reduced during
40 drought, we can expect to see a shift towards more resilient methanogens and methanotrophs under
41 future climate conditions.

42 Plain Language Summary

43
44 Droughts and heat waves are increasing due to climate change. This can lead to considerable stress
45 on soil microorganisms in northern wetlands which emit strong greenhouse gases such as methane.
46 A better understanding of these methane producing and consuming microorganisms can help us
47 predict the how the community responds to droughts, and thus, how much green house gases will
48 be emitted in the future contributing further to climate change.

50 1 Introduction

51 Anthropogenic climate change is one of the key issues of the twenty-first century, and it has the
52 potential to severely impact natural peatlands through changes in temperature and precipitation
53 (IPCC, 2021). While climate change models are forecasting increased precipitation at northern
54 latitudes, these events are predicted to be more concentrated and less frequent in time with longer
55 periods of dryer warm weather in between (IPCC, 2021). These events often result in a lowering
56 of the water table depth which exposes methanogenic anaerobes to oxidative stress. This may
57 decrease methane (CH₄) emissions to the atmosphere by reducing the habitable anoxic zone where
58 methanogenic Archaea produce CH₄, but potentially also by increased activity and abundance of
59 methanotrophs leading to a higher consumption of CH₄ (Keane et al., 2021, Rinne et al., 2020).
60 These microbial communities inhabiting natural peatlands are vulnerable to disturbance under a
61 warming climate, but potential structural shifts in microbial communities are currently difficult to
62 predict, contributing to high uncertainties in current CH₄ budgets (Dean et al., 2018, Saunio et al.,
63 2020).

64
65 Pristine peatlands function as long-term carbon sinks because plant productivity (CO₂ uptake)
66 generally exceeds the slow rate of organic matter decomposition (CO₂ release) due to anaerobic

67 conditions. Although these anaerobic conditions lead to significant emissions of CH₄, over long-
68 time scales, the carbon balance of peatlands is primarily determined by the CO₂ fluxes (Yu, 2012,
69 Rinne et al., 2020). However, the greenhouse gas balance of a peatland can shift from a sink to a
70 source after drought, when expressed in CO₂-equivalents, due to the higher global warming
71 potential of CH₄ (Fenner and Freeman, 2011, Rinne et al., 2020).

72
73 Drought conditions, i.e. high air temperatures and reduced precipitation, result in a lower water
74 table depth, aerating previously anoxic peat layers. This leads to increased heterotrophic
75 respiration, and consequently a higher release of CO₂ to the atmosphere (Keane et al., 2021, Rinne
76 et al., 2020). Concurrently, CH₄ emissions are reduced since oxygen (O₂) inhibits CH₄ production
77 upon exposure to methanogen cells combined with increased methanotrophic CH₄ oxidization
78 (Miller et al., 2019, Thauer et al., 2008). Due to these microbial controls, a deeper understanding
79 of microbial structure and function, and the relationship to hydrological status, is required to
80 improve projections of the role of peatland GHG emissions in the climate system.

81
82 Under anoxic conditions, inhibitory phenolic compounds are built up. These compounds prevent
83 the activity of polyphenolic carbon degrading aerobes, which enable greater conversion of peat
84 organic carbon into smaller substrates such as sugars, organic acids, H₂, and CO₂ that are more
85 bioavailable for the anaerobic methanogens (Wilmoth et al., 2021, Fenner and Freeman, 2011).
86 However, if O₂ is introduced through a drop in water table depth, phenol oxidase can remove
87 phenolic inhibitors, enabling hydrolases to resume normal mineralization of organic matter that
88 subsequently provide additional substrates for methanogenesis upon the return to anoxic conditions
89 (Fenner and Freeman, 2011, Wilmoth et al., 2021).

90
91 The taxonomic structure and function of methanogens is diverse and closely linked to hydrology
92 status and warming (Bräuer et al., 2020). CH₄ production occurs stepwise in cooperation between
93 different microbial functional groups, where organic carbon bound to dead organic matter is
94 converted into CH₄ via methanogenesis (Ferry, 1999, Dean et al., 2018). Methanogenesis is a
95 process catalyzed by specialized functional groups that convert CO₂ with H₂, methanol,
96 methylamines, methylsulfides, or acetate into CH₄ (Thauer et al., 2008). Anaerobic
97 methanogenesis is carried out exclusively by members of the archaeal domain (Bräuer et al., 2020).
98 Methanogens display high phylogenetic diversity spanning three phyla (*Euryarchaeota*,
99 *Halobacterota* and *Thermoplasmata*) and are no longer considered strict *Euryarchaeota*
100 members. In total, five orders and two candidate taxa are commonly discovered in peat:
101 *Methanomicrobiales*, *Methanocellales* and *Methanosarcinales* of the phylum *Halobacterota*;
102 *Methanobacteriales* of the phylum *Euryarchaeota*; *Methanomassiliicoccales* of the phylum
103 *Thermoplasmata*, and finally candidate family *Methanoflorentaceae* and candidate phylum
104 *Bathyarchaeota* (Bräuer et al., 2020).

105
106 In contrast to methanogens, methanotrophs – of the phyla, *Proteobacteria*, *Verrucomicrobia*, and
107 candidate phylum NC10 – can oxidize CH₄ before it is emitted to the atmosphere, acting as a
108 natural bio-filter. Methanotrophs commonly inhabit the oxic-anoxic interfaces, where they oxidize
109 between 10 to 90% of the CH₄ produced by methanogens (Hakobyan and Liesack, 2020,
110 Wendlandt et al., 2010).

111
112 In this study, we focus on the drought that occurred during the summer of 2018 when Northwestern

113 Europe, including Sweden, experienced a heatwave (Rinne et al., 2020, Sjökvist et al., 2019,
114 Vicente-Serrano et al., 2010). Lower precipitation and higher temperatures altered the hydrological
115 status of Swedish peatlands, leading to a lower water table depth, increased peat temperatures, and
116 altered biogeochemical processes – including changes to methanogenesis. Although,
117 methanogenesis is one of the most important carbon degradation pathways in peatlands (Keane et
118 al., 2021, Kelly et al., 2021), knowledge on the resilience of methanogenic archaea to droughts in
119 terms of community abundance, diversity and structure is still poorly understood and requires
120 further attention (Kim et al., 2008).

121
122 Here, we address the functional potential of methanogenic and methanotrophic microbes in
123 response to drought. We hypothesise that (1) the proportion of methanogenic and methanotrophic
124 community shifts in relative abundance towards higher methanotrophic abundances when exposed
125 to drought conditions. In addition, we aim to (2) determine which vegetative ecotype holds the
126 highest microbial diversity during the drought and (3), to identify whether the functional gene
127 composition shifts in response to drought.

128

129 **2 Materials and Methods**

130

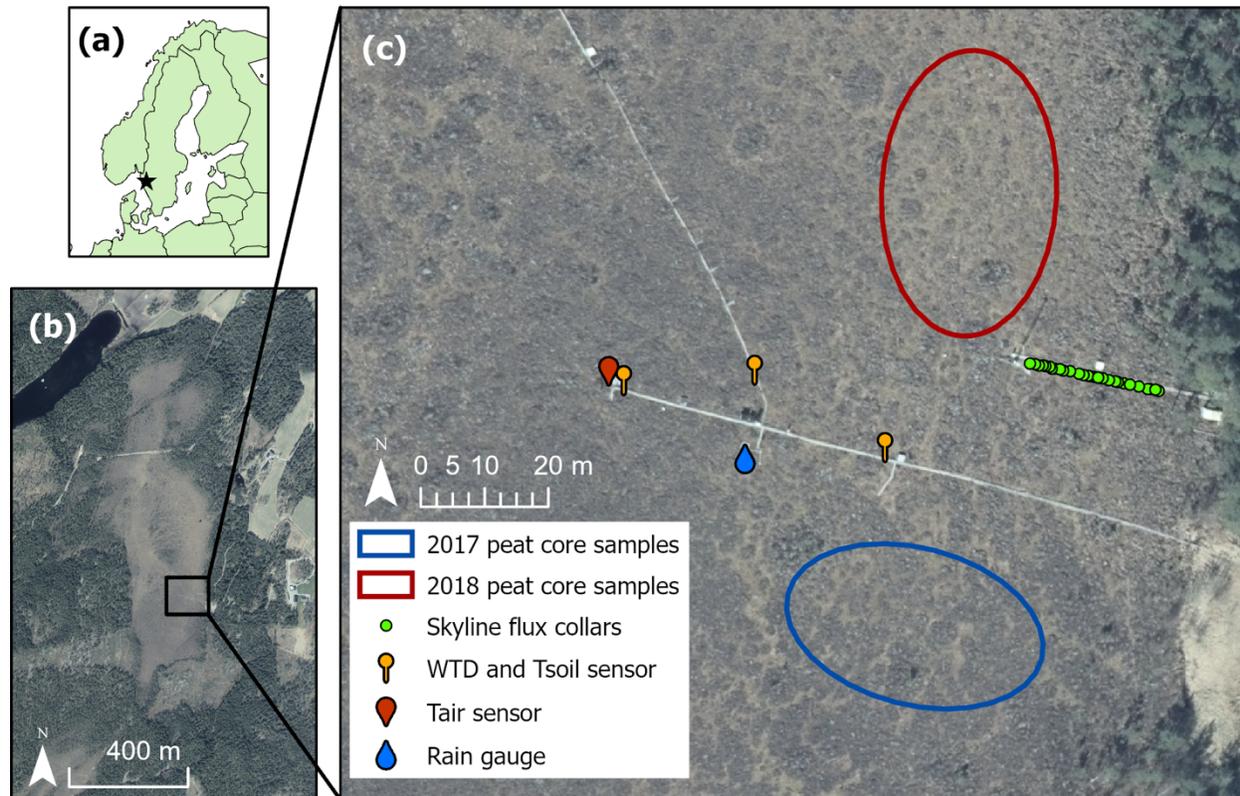
131 **2.1 Site Description**

132 This study focuses on Mycklemossen, a hemi-boreal mire dominated by bog-like vegetation,
133 located in southern Sweden (58°21'N, 12°10'E). Mycklemossen is a sub-section of the Skogaryd
134 Research Catchment and Swedish Infrastructure for Ecosystem Science network
135 (<https://www.fieldsites.se>). Common to many hemi-boreal mires, the peatland consists of wet low
136 areas dominated by *Sphagnum rubellum* and *Rhynchospora alba*, while the raised intermediate
137 areas are a result of the tussock-building sedge *Eriophorum vaginatum*. Once the tussocks are
138 established, the upper layers of the peat become drier and are no longer anoxic. This allows for the
139 establishment of low shrubs such as *Calluna vulgaris*. The long-term (1990–2019) mean annual
140 air temperature and total precipitation were 6.7°C and 1021 mm respectively, as measured by the
141 closest national monitoring station (Vänernsborg, 10 km to the east and at a 30 m lower elevation
142 than Mycklemossen).

143 **2.2 Experimental design**

144 To determine the impact of the 2018 drought, we measured CH₄ fluxes, soil and air temperature,
145 precipitation, water table depth, and collected peat samples for genetic analysis during 2017 and
146 2018. Measurements of CH₄ were made across a ~40 m long transect beginning at the tree line and
147 extending into the mire. Plots were classified according to the dominant vegetation and are here
148 represented as the *E. vaginatum* (n = 6), *C. vulgaris* (n = 6) and *R. alba* (n = 6) ecotypes. Replicate
149 plots were identified at random and classified according to the dominant vegetation type. Peat
150 samples for extraction of gDNA were removed from two different locations north and south of the
151 boardwalk displayed in Figure 1. In 2017, 18 peat samples were collected from locations
152 representing the *C. vulgaris* (n = 6), *E. vaginatum* (n = 8), and *R. alba* (n = 4) ecotypes south of
153 the boardwalk (Figure 1). In 2018, 11 samples were collected from *C. vulgaris* (n = 2), *E.*
154 *vaginatum* (n = 5), *R. alba* (n = 4) ecotypes north of the boardwalk (Figure 1). Peat sampling was
155 conducted in two similar sampling locations north and south of the boardwalk and these were

156 separated by a distance of ~60m to avoid disturbance from previous sampling events (Figure 1).
 157 Both areas are similarly composed of hummocks and hollows and include equal representation of
 158 the pre-described ecotypes, which each exist within their own hydrological niche. The majority of
 159 the local peat deposits within the catchment extend down to 6 m depth (Wallin et al., 2015), with
 160 previous studies establishing no significant differences in surface or soil temperature between the
 161 hummocks and hollows (Kelly et al., 2021).



162
 163 **Figure 1:** (a) Location of Mycklemossen (black star) within Sweden (b) aerial photo of
 164 Mycklemossen, black square shows the location of the sampling area and (c), map of peat core and
 165 ancillary measurement locations. Map data sources: © EuroGeographics and © Lantmäteriet.

166

167 2.3 Environmental variables

168 Meteorological variables were measured in both 2017 and 2018 with which we characterized the
 169 severity of the drought at Mycklemossen. Air temperature was measured with an HC2S3 sensor
 170 (Campbell Scientific, Logan, UT, USA) at 2 m above the peatland surface in a ventilated, radiation
 171 protected housing. Precipitation was measured with a tipping bucket rain gauge (SBS500H,
 172 Campbell Scientific, United Kingdom). Water table depth (CS450, Campbell Scientific, UT United
 173 States) was measured at three different locations that represented the *C. vulgaris*, *E. vaginatum* and
 174 *R. alba* ecotypes. The locations for soil and air temperature, precipitation and water table depth
 175 sensors are shown in Figure 1. All environmental variables were measured at 1 Hz and recorded
 176 on a CR1000 data logger (Campbell Scientific, UT United States). Air temperature and water table
 177 depth values were averaged to represent daily means, while precipitation was summed to represent
 178 total daily values.

179 2.4 CH₄ flux measurements

180 Surface GHG fluxes were measured using the SkyLine2D system. The SkyLine2D is an automated
181 chamber system designed and built at the University of York to measure greenhouse gas exchange.
182 For a full description of the SkyLine2D system, we refer to Keane et al. (2018). In short, the flux
183 chamber comprised of a translucent Perplex cylindrical chamber (inner diameter 20 cm, height 40
184 cm), which was suspended from a motorized trolley and programmed to traverse ca. 2 m above the
185 transect. The system was preset to visit the pre-selected plots along the transect, where the chamber
186 was lowered onto pre-installed collars for a measurement period of 4 minutes. Following the 4-
187 minute measurement period, the system raised the chamber and moved to the next plot. The time
188 taken to complete a full cycle was approximately 2.5 h, which allowed each chamber to be
189 measured ca. 10 times per day. The headspace gas from within the sealed chamber was circulated
190 through a Los Gatos cavity ring-down laser (CRD, LGR U-GGA-91, Los Gatos Research, CA
191 United States) to measure the change in concentration of CH₄. Fluxes were calculated as the
192 increase in headspace concentration over time, determined by linear regression, and adjusted for
193 temperature and area of the chamber.

194

195 2.5 Captured metagenomics

196 2.5.1 DNA extraction

197 Peat material was collected using a 1.5m long box corer from different locations within the mire
198 that best represented the dominant ecotypes (Figure 1). Small samples of peat (~30 grams) were
199 collected from the oxic-anoxic interface at ~5cm and within the anoxic zone, at ~30cm. Once
200 separated from the cores, the peat material was immediately snap frozen using liquid nitrogen and
201 stored in a -20°C freezer. Before DNA extraction was performed, samples were thawed in a
202 refrigerator at 4°C. After thawing, gDNA was extracted from peat samples following the DNeasy®
203 PowerSoil® Kit (Qiagen, Hilden, Germany) and carried out according to the manufacturers
204 protocol, including the recommended 0.25g of input material. Following the DNA extraction,
205 samples were tested for quality (absorbance ratio 260/280) and concentration on a NanoDrop lite
206 (NanoDrop Technologies, Wilmington NC, USA) and Invitrogen Qubit 4 fluorometer (Thermo
207 Fisher Scientific, Waltham MA, USA) respectively.

208

209 2.5.2 SeqCap EZ probe generation

210 The metagenomic DNA extracted from the peat was processed to enrich for sequences of interest
211 via the “captured metagenomics” method using oligonucleotide probes following White et al.
212 (2022) and Manoharan et al. (2015). In short, genes that encode enzymes related to methane
213 production and consumption were identified from the Kyoto Encyclopedia of Genes and Genomes
214 database (KEGG) (Kanehisa et al., 2015). In total, 548,104 genes coding for methane metabolism
215 were downloaded via a custom R script (<https://github.com/dagahren/metagenomic-project>) and
216 compiled into a local database, subsequently called the CH₄ database. The nucleotide coding
217 sequences of the CH₄ database were used to design custom hybridisation-based probes for
218 sequence capture according to Kushwaha et al. (2015). In total, 193,386 individual probes were
219 generated after clustering, with a melting temperature of 55°C and probe length 40mer, suitable
220 for use with the NimbleGen SeqCap EZ protocol (Roche NimbleGen Inc., Madison, USA).

221

222 *2.5.3 Library generation, probe hybridisation and sequencing*

223 Depending on the concentration of the extracted DNA in a total volume of 100µl low TE, either
224 150ng or 1µg of gDNA was sheared using a Bioruptor Pico in 0.65ml Bioruptor tubes for 13 cycles
225 – 30s on, 30s off (Diagenode SA, Seraing, Belgium). The fragmented DNA was purified using
226 1.8× AMPure XP beads (Beckman Coulter) and used as input material for preparation of pre-
227 capture libraries. Libraries were constructed according to the Nimblegen SeqCap EZ HyperCap
228 Workflow User's Guide (Version 1.0, June 2016) with the following modifications: (1) for the
229 adapter ligation step, 5µl of 15µM KAPA unique dual index mixed adapters were used instead of
230 single index adapters, (2) for the pre-capture PCR, 7 cycles were used for libraries with a genomic
231 DNA input of 150ng, and 5 cycles where the input was 1µg.

232

233 Libraries were multiplexed in pools of 15 in equimolar amounts based on the concentrations and
234 sizes. 1µg of each pool was transferred to a test tube and hybridised to the custom probes according
235 to the NimbleGen SeqCap EZ SR User's Guide (Version 4.3, October 2014). The capture tubes
236 were incubated in a thermal cycler set at 47 °C, heated lid set to 57 °C for 69 hours. The quantity
237 and quality of the final pool was assessed by Qubit and Bioanalyzer and subsequently by qPCR
238 using the Illumina Library Quantification Kit from Kapa on a Roche Light Cycler (LC480II, Basel,
239 Switzerland).

240

241 The captured libraries were sequenced on an Illumina HiSeq4000 platform using sequencing by
242 synthesis technology to generate 2 x 150 base paired end reads. The analysis was carried out at the
243 Centre for Genomic Research, University of Liverpool, United Kingdom.

244

245 **2.6 Data processing and statistics**246 *2.6.1 Environmental variables*

247 All environmental data, including soil and air temperatures, precipitation and water table depths
248 were measured from the 1st of May to the 30th of September, which we refer to as the growing
249 season. In addition, we tested for differences in the mean values of air temperature, soil temperature
250 and water table depth, according to the defined ecotypes, between 2017 and 2018.

251

252 *2.6.2 CH₄ flux*

253 CH₄ flux data was quality controlled by discarding measurements with a R² value ≤ 0.9.
254 Measurements passing this threshold were then assessed using the output statistics from the
255 regression calculation, where regressions with a p value ≤ 0.05 were accepted, while those that did
256 not were treated as zero flux. To allow for the temporal and repeated measures data, differences in
257 CH₄ fluxes between years and ecotypes were tested using linear mixed effects models via the lme4
258 package v1.1-27.1 (Bates et al., 2015). Differences were calculated using estimated marginal
259 means in combination with a Tukey pairwise post-hoc tests on significant effects.

260

261 *2.6.3 Sequence annotation*

262 Raw sequencing files were trimmed for the presence of Illumina adapter sequences using the
263 software package Cutadapt v1.2.1 (Martin, 2011). The reads were further trimmed

264 using Sickle v1.2 with a minimum window quality score of 20 (Joshi, 2011). Following trimming,
265 reads shorter than 20bp were discarded. The sequence reads from each of these captured data sets
266 were processed through MG-RAST, an online metagenomics annotation program (Meyer et al.,
267 2008). Default parameters were used for quality filtering of low quality reads and removal of
268 sequence duplicates. The taxonomic and functional annotations from MG-RAST were annotated
269 using the RefSeq (O'Leary et al., 2016) and KEGG (KO) (Kanehisa et al., 2015) databases.
270 Following the MG-RAST pipeline, sequences were further filtered for both taxonomic and
271 functional gene annotations using the KEGG methane metabolism filter (ko:00680). The filter
272 includes both taxonomic and functional genes related to methane metabolism and excludes
273 remaining off target sequences. Sequence data and functional annotations are freely available
274 through MG-RAST with the accession ID mgp91145. These filtered sequences were then exported
275 to R for further analysis.
276

277 2.6.3 Taxonomic diversity and functional genes

278 Diversity indices were calculated via the phyloseq package v1.3.0 (McMurdie and Holmes, 2013)
279 where taxonomic abundances below 10 reads were removed. Due to the small sample sizes and
280 uneven distribution of replicates, a PERMANOVA was used with 999 permutations (Anderson,
281 2001). First, we normalized taxonomic and functional gene relative abundance via a double square
282 root transformation to allow for highly abundant genes. Following transformation, we calculated
283 ordination using Bray-Curtis distances and finally, a Wilkson pairwise post-hoc test was used to
284 identify significant differences between ecotypes and years with the vegan package v2.5 (Oksanen
285 et al., 2019). All analyses were completed in the R statistics package v 3.6.1 (R Core Team, 2018)
286 and visualized using the ggplot2 package v 3.3.2 (Villanueva and Chen, 2019).
287

288 3.0 Results

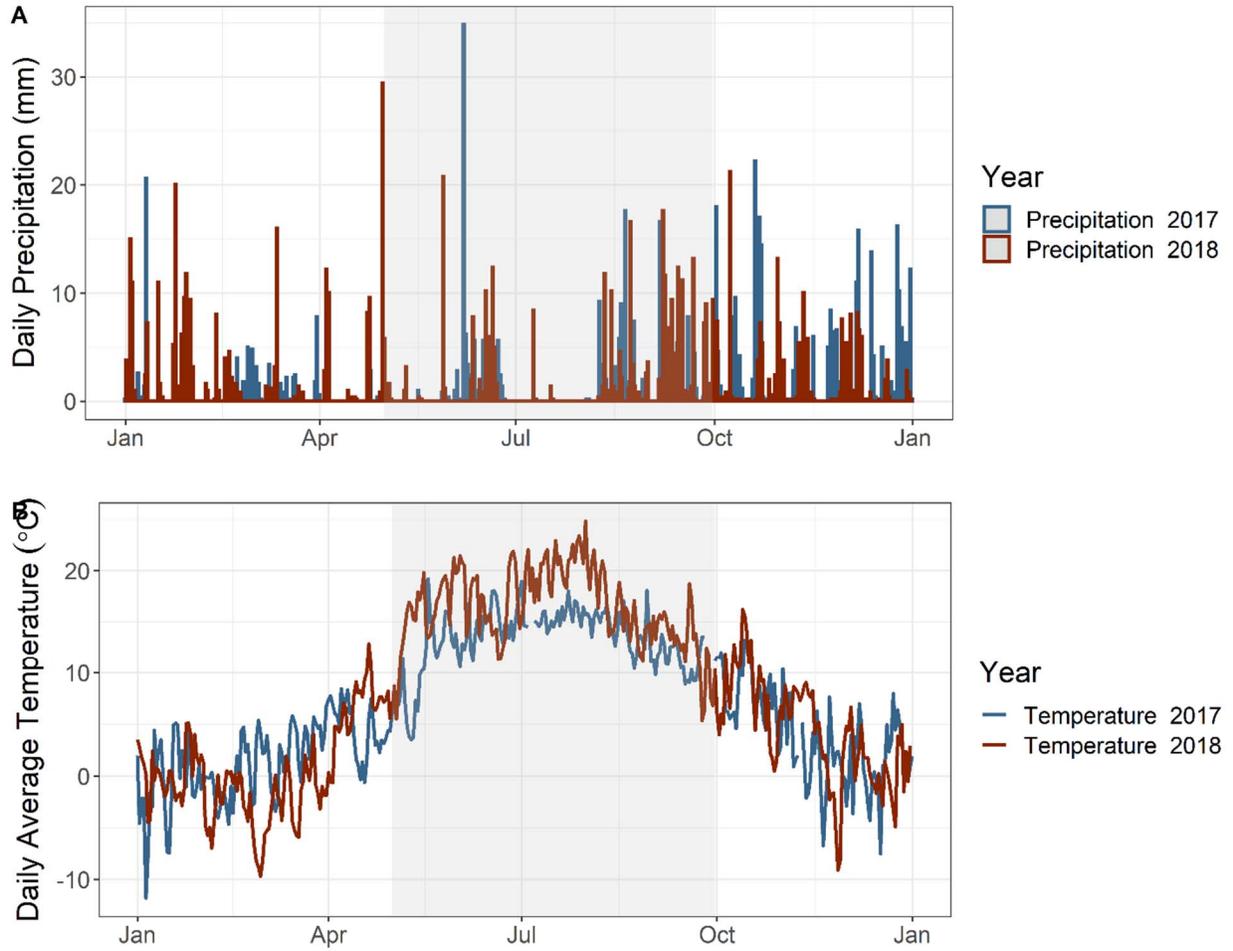
289

290 3.1 Environmental variables

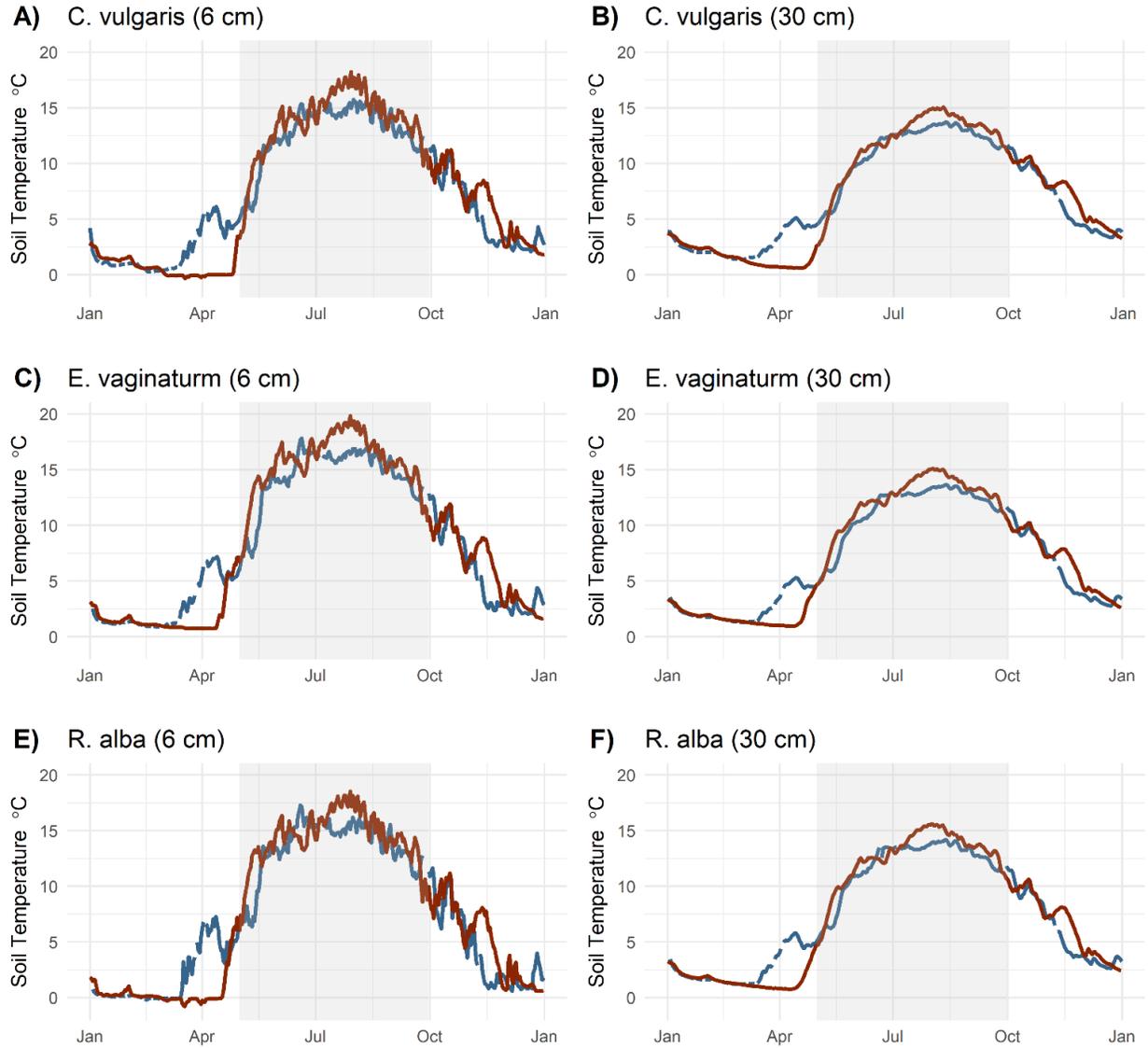
291 In 2018, Mycklemossen mire experienced a maximum air temperature of 33°C that was reached
292 on the 31st of July 2018 (Figure 2A). Daily mean air temperature during May-September reached
293 a maximum of 24.8°C (SD ± 3.95) in 2018, a 5.6°C increase in comparison to 2017. On average,
294 all daily mean soil temperatures during the growing season were higher in 2018 when compared
295 to 2017. The high air temperature was also reflected in the soil temperature, where the maximum
296 daily values in August 2018 were 3.4°C higher at 6cm depth and 1.3°C higher at 30cm depth in
297 the *C. vulgaris* ecotype compared to 2017 values (Figure 3, A and B). The *E. vaginatum* ecotype
298 followed the same pattern with 2°C higher at 6cm depth and 1.5°C higher at 30cm depth (Figure
299 3, C and D). The final ecotype, *R. alba*, had a slightly lower deviation with 1.4°C higher at 6cm
300 depth and 1.2°C higher at 30cm depth (Figure 3, E and F).
301

302 Daily summed precipitation values in 2018 were below average during the growing season from
303 May to October (Figure 2B), with predominantly dry conditions in July 2018 when only 10.2 mm
304 of rain was recorded, compared to the long-term average of 80 mm. Significantly lower rainfall
305 was observed in 2018 during the whole growing season when compared to 2017 values ($p \leq 0.04$)
306 (Figure 3). This decrease in precipitation resulted in a lower water table depth across the whole
307 peatland (Figure 4). Significantly lower WTD's were observed in both *E. vaginatum* ($p \leq 0.02$)
308 and *R. alba* ($p \leq 0.02$), resulting in a lower water table depth up to 4 cm and 9 cm in 2018 (Figure

309 4B). In the *C. vulgaris* ecotype, the water table dropped by as much as 7 cm in 2018 when compared
 310 to 2017.
 311

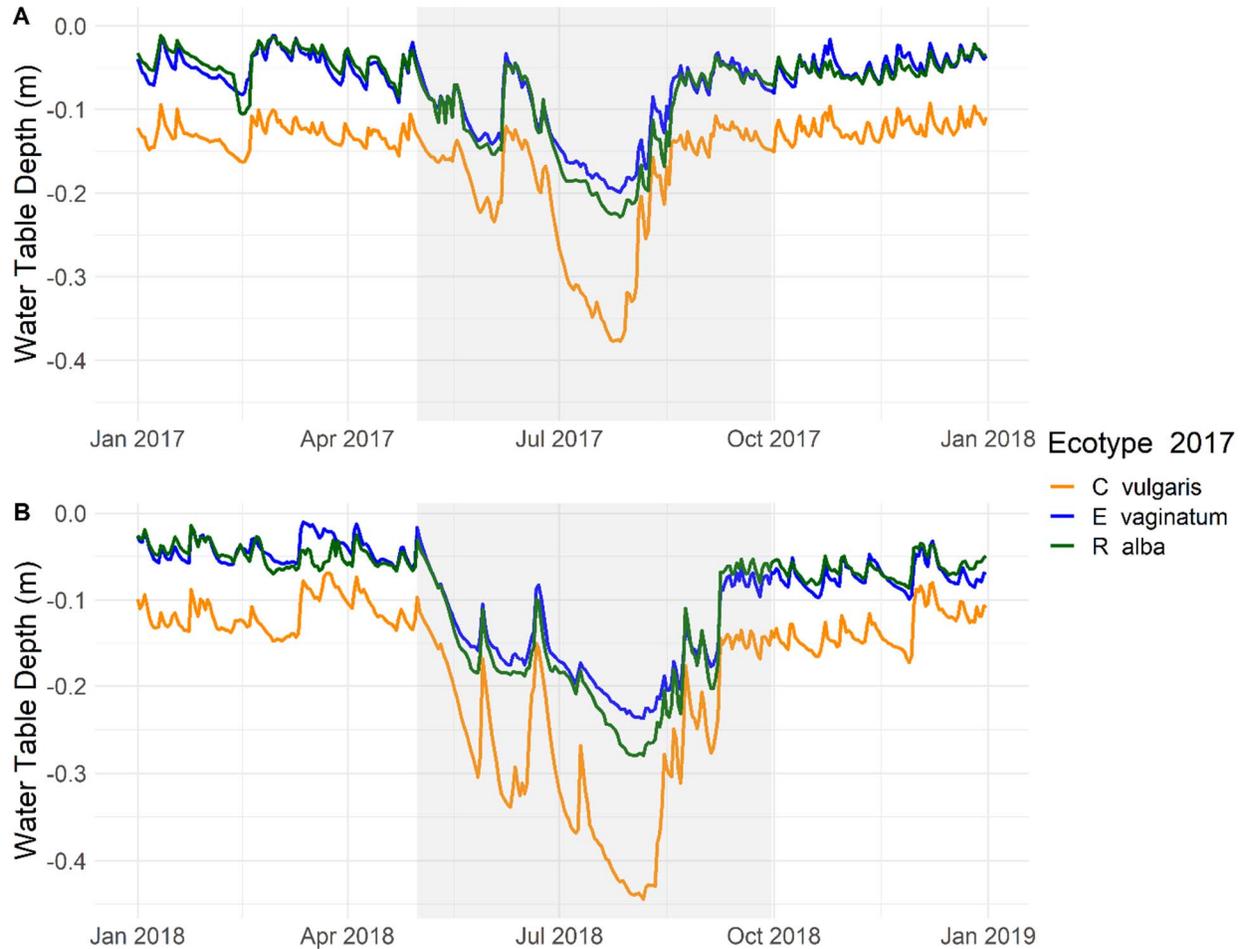


312 **Figure 2:** A) Daily mean temperature measured in °C and B) Daily summed precipitation
 313 measured in millimeters at Mycklemossen mire. 2017 is shown in blue and 2018 in red. The shaded
 314 area indicates the growing season from May 1st to September 30th.
 315



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Figure 3: Daily mean soil temperature measured at 6 cm and 30 cm below the peat surface. Blue lines indicate 2017 values while red lines indicate 2018 values. The shaded area indicates the growing season from May 1st to September 30th.

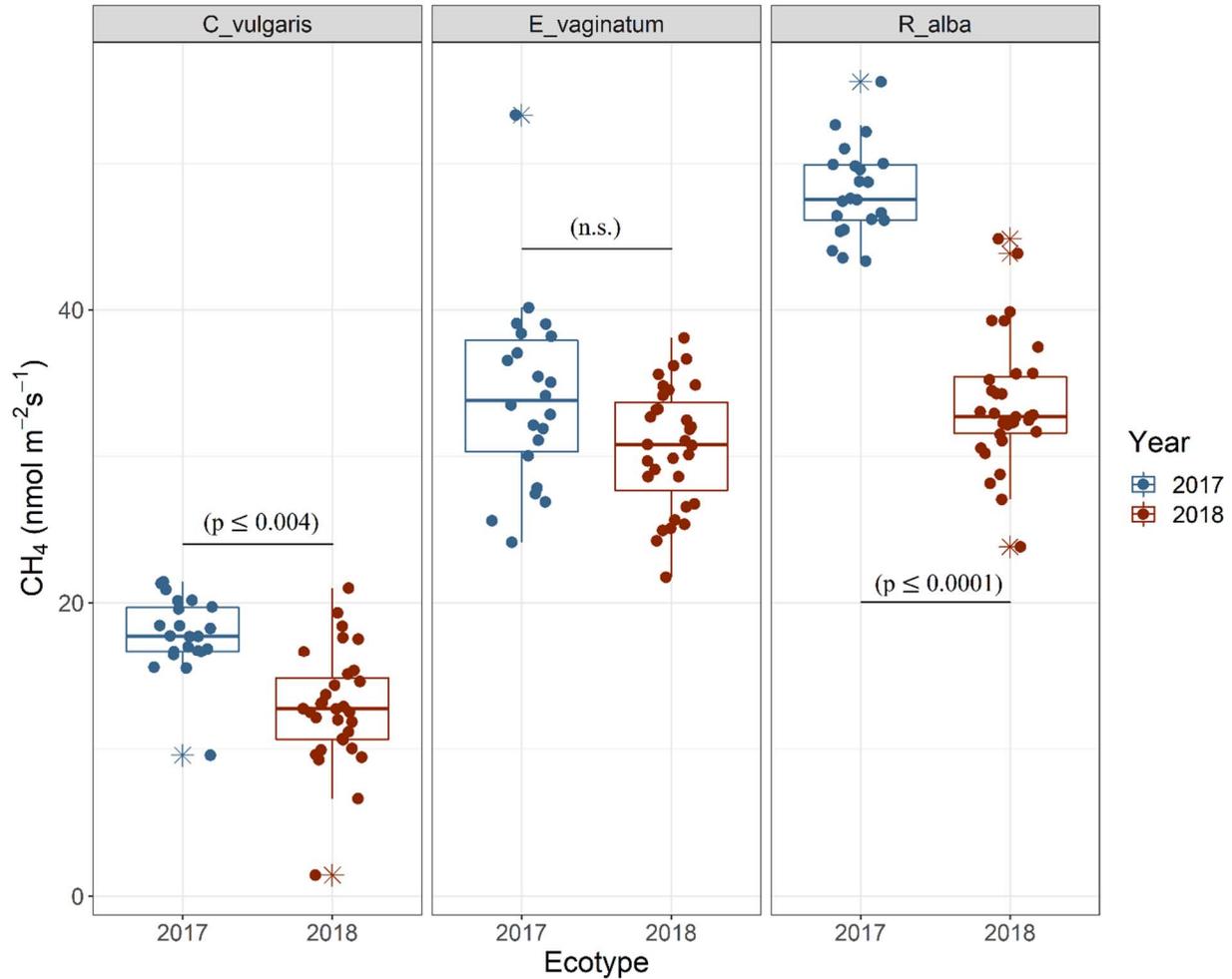


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Figure 4: Water table depth in meters below the surface for *C. vulgaris*, *E. vaginatum* and *R. alba* ecotypes from automated sensors at the site during 2017 (A) and 2018 (B). The shaded area indicates the growing season from May 1st and September 30th.

325 3.2 Methane fluxes

326 The highest daily mean CH₄ fluxes (33.3 CH₄ nmol m⁻² s⁻¹) were observed during the growing
327 season in 2017, significantly higher than in 2018 (25.7 CH₄ nmol m⁻² s⁻¹) ($p \leq 0.001$). The ecotype
328 that yielded the highest mean CH₄ flux was *R. alba* in 2017 with a mean flux of 48.1 CH₄ nmol m⁻²
329 s⁻¹ (SD \pm 3.11) followed by *E. vaginatum* (34.1 SD \pm 6.36 CH₄ nmol m⁻² s⁻¹) and *C. vulgaris*
330 (17.8 SD \pm 2.58 CH₄ nmol m⁻² s⁻¹). During the drought in 2018, fluxes in the *R. alba* ecotype
331 significantly reduced by 29% ($p \leq 0.0001$), and by 27% ($p \leq 0.04$) in the *C. vulgaris* ecotype, while
332 the smallest reduction was observed in *E. vaginatum* (10%) ($p \geq 0.05$).
333



334
 335 **Figure 5:** Boxplots of daily mean CH₄ flux measured during the growing season (May to October)
 336 in 2017 (blue) and 2018 (red) at Mycklemossen mire. The boxes show quartiles and the median,
 337 the whiskers denote data within 1.5 times of the interquartile range. Colored stars denote outliers
 338 while significant differences using linear mixed effects models are labelled with the p value and
 339 “n.s.” indicates a non-significant result.

340
 341 **3.3 Taxonomy**

342 *3.3.1 Proportions of methanogens to methanotrophs between years*

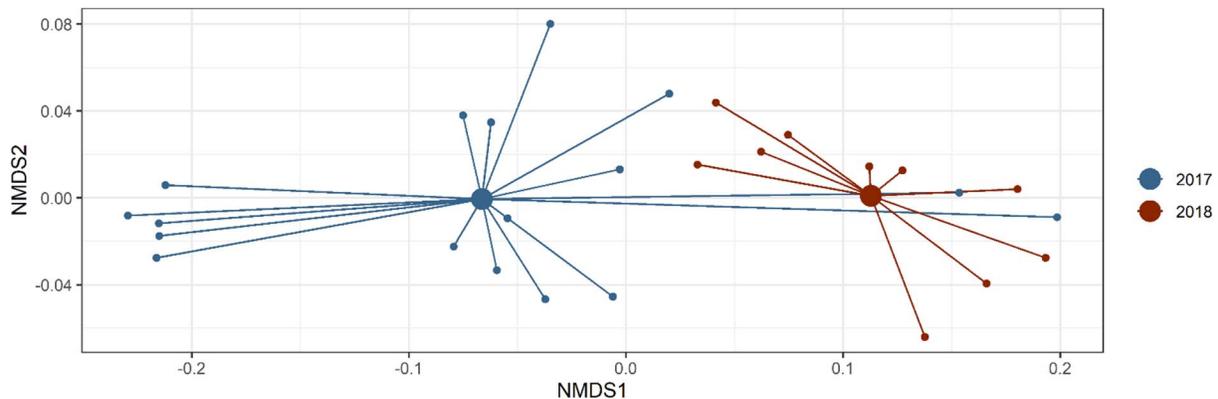
343 Large significant variations in the abundances of methanogens and methanotrophs were observed
 344 between drought and non-drought years, with 36% of the variation explained by the year ($R^2 =$
 345 0.36, $p \leq 0.003$) (Figure 6). Clear clusters can be observed in Figure 6 with only a small overlap in
 346 samples. The average proportion of methanogens to methanotrophs during 2017 was 59% and 41%
 347 respectively. In 2018, however, the proportion of methanogens decreased by 8%, while the
 348 proportion of methanotrophs increased to 49%. The genus which contributed the most to
 349 dissimilarity according to the SIMPER analysis – that calculates the contribution of each species
 350 (%) to the dissimilarity between each group – was *Methylocella*, with an 0.12 average contribution
 351 to the overall dissimilarity (Table 1). The average abundance for *Methylocella* significantly
 352 increased by 219% during the drought in 2018 when compared to 2017 ($p \leq 0.001$). Interestingly,

353 the average abundance of all methanogens and methanotrophs increased during the drought, with
 354 *Methanoregula* contributing the highest to dissimilarity, followed by *Methanosarcina*. In addition,
 355 all detected methanotrophs significantly increased, including the type I, II and *Verrucomicrobia*
 356 when tested between years ($p \leq 0.005$).

357
 358 **Table 1:** Results of taxonomic contribution of each species (%) to the dissimilarity between each
 359 group (SIMPER analysis) between 2017 and 2018. Taxa are ranked according to their average
 360 contribution to dissimilarity between years. Average abundances, percentage of cumulative
 361 contribution and Permutation *p*-value (Probability of getting a larger or equal average contribution
 362 in random permutation of the group factor) are also included. A cut-off at a cumulative
 363 dissimilarity of 70% was applied (2017 $n = 17$; 2018 $n = 10$).

Genus	Average contribution to dissimilarity	SD	Avg. 2017	Avg. 2018	Percentage contribution	p-value
<i>Methylocella</i>	0.12	0.089	2539	8108	20%	0.001
<i>Methanoregula</i>	0.09	0.085	4267	5968	36%	1
<i>Methylosinus</i>	0.08	0.050	2386	5602	49%	0.031
<i>Methanosarcina</i>	0.05	0.032	1529	4142	58%	0.006
<i>Methylococcus</i>	0.03	0.022	814	2495	64%	0.001
<i>Methylacidiphilum</i>	0.03	0.020	568	1935	69%	0.001

364



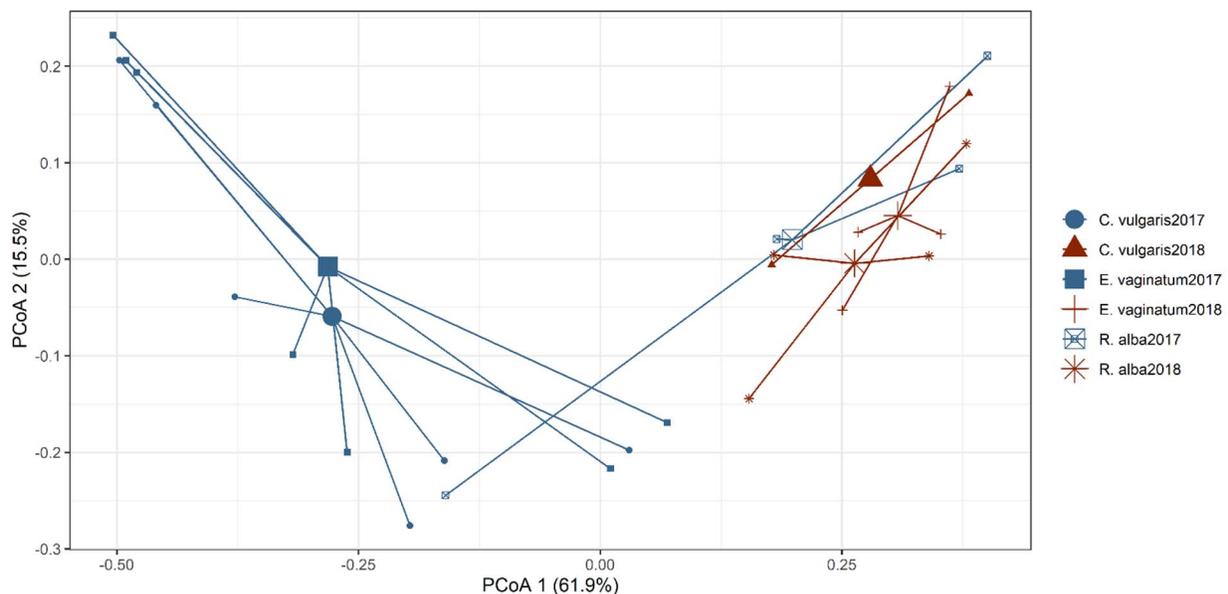
365
 366 **Figure 6:** Non-metric Multidimensional Scaling (NMDS) of taxonomic abundances using Bray-
 367 Curtis distances. Small dots indicate individual samples while the largest dots indicate the mean of
 368 all samples. Samples were analyzed at genus level and colored by year 2017 ($n = 17$) and 2018 (n
 369 $= 10$).
 370

371 3.3.2 Ecotype comparison

372 A large dissimilarity in the relative abundance of taxa was observed when comparing ecotypes
 373 between each year. This variation in taxonomic abundance resulted in a significant correlation
 374 where 58% of the variation in taxonomic abundance could be explained by ecotype ($R^2 = 0.58$, p
 375 ≤ 0.002). The Wilks pairwise test indicated that significant dissimilarity occurred for *E.*
 376 *vaginatum* between 2017 and 2018 ($p \leq 0.03$), but the same dissimilarity was not observed for *R.*
 377 *alba* and *C. vulgaris* ($p \geq 0.05$).

378 When comparing the relative abundance of methanogens and methanotrophs between *E.*
 379 *vaginatum*, *R. alba* and *C. vulgaris*, six genera were the most common within the top 70% of
 380 cumulative sums, irrespective of ecotype: *Methanoregula*, *Methylocella*, *Methylosinus*,
 381 *Methanosarcina*, *Methylococcus* and *Methylacidiphilum* (Supplementary table 1, 2 and 3).
 382 *Methanocaldococcus* and *Methanosphaerula* were within the top 70% of cumulative sums for the
 383 *R. alba* ecotype (Supplementary table 2), but not *E. vaginatum* (Supplementary table 1) or *C.*
 384 *vulgaris* (Supplementary table 3).

385 When comparing *R. alba* and *C. vulgaris*, we observe that the hydrogenotrophic *Methanoregula*
 386 contributed the most to dissimilarity (Supplementary Table 2 and 3). Although type II
 387 *Alphaproteobacteria* genera, *Methylocella* and *Methylosinus* were the second and third most
 388 important between ecotypes, the order of dissimilarity of *Methylocella* and *Methylosinus* changed
 389 according to ecotype. Within the *E. vaginatum* comparison, the dominant *Methanoregula* was
 390 surpassed by *Methanosarcina*, with an average dissimilarity of 0.07, when compared to the
 391 remaining ecotypes.



392

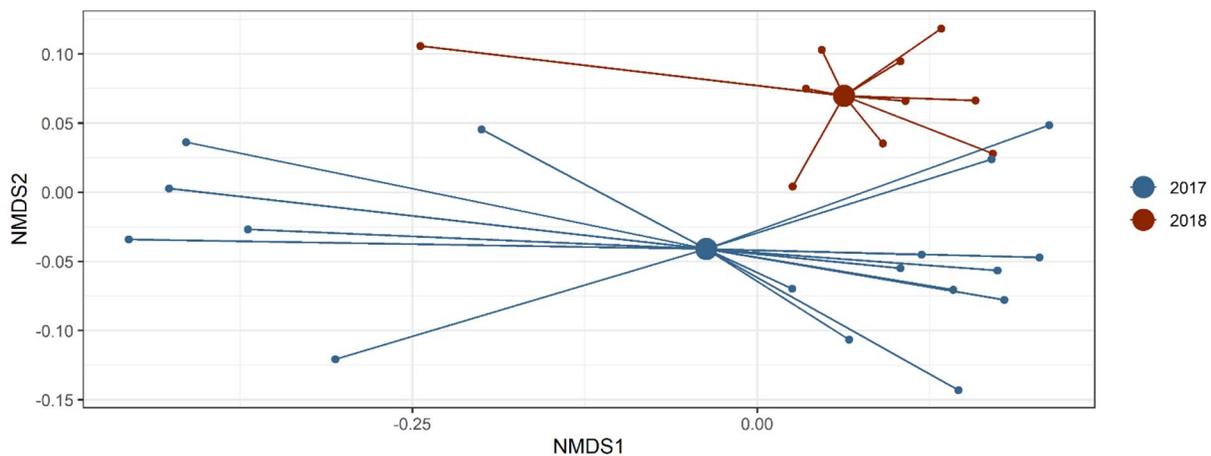
393 **Figure 7:** Principal Coordinates Analysis (PCoA) of taxonomic abundances using Bray-Curtis
 394 distances. Small symbols indicate individual samples while the largest symbols indicate the mean
 395 of all samples. Samples were analyzed at genus level and colored by ecotype and sampling year

396 (*C. vulgaris* 2017 (n = 6) 2018 (n = 2), *E. vaginatum* 2017 (n = 7) 2018 (n = 4), *R. alba* 2017 (n =
 397 4) 2018 (n = 4).

398

399 3.5 Functional gene composition

400 The functional genes showed a clear separation between 2017 and 2018, with only a small overlap
 401 between clusters (Figure 8). The largest variation of functional genes occurred in 2017, with
 402 smaller variation observed in 2018. Significant differences were observed between the abundance
 403 of functional genes when tested via PERMANOVA between 2017 and 2018 ($p \leq 0.036$). In
 404 addition, the PERMANOVA revealed that 12% of the variance in abundances can be explained by
 405 the year ($R^2 = 0.12$, $p \leq 0.036$).



406

407 **Figure 8:** Nonmetric Multidimensional Scaling (NMDS) of functional genes using Bray-Curtis
 408 distances. Small dots indicate individual samples while the largest dots indicate the mean of all
 409 samples. Samples were analyzed at KO level 4 and colored by year 2017 (n = 17) and 2018 (n =
 410 10).

411

412 In total, 106 functional genes related to CH₄ metabolism were captured, with 20 contributing to
 413 the top 70% cumulative sum of dissimilarity between CH₄ functional genes (Supplementary Table
 414 4). Within the top 70% cumulative contributions, 12 out of the 20 captured genes saw an increase
 415 in average abundance in 2018. The gene contributing most to the dissimilarity was heterodisulfide
 416 reductase subunit A (*hdrA*). The *hdrA* gene held an average abundance of 627 in 2017, resulting
 417 in an 86% increase in 2018 ($p \geq 0.05$). Genes including *cutL*, *hdr*, *fdhA*, *coxS*, *frmB*, *mvhA*, *metF*
 418 and *cutM* were all significantly more abundant in 2018 when compared to 2017 values ($p \leq 0.05$).
 419 Genes that did not increase during 2018 included *frhA*, *mcrA*, *frhG*, *hdrB*, *fwdB*, *mtd*, *mtrE* and
 420 *mtrH*.

421

422 4.0 Discussion

423

424 In this study, we observed the effect of drought on the functional potential of CH₄ producing and
 425 reducing microorganisms. In general, methanogens and methanotrophs displayed a high resilience

426 to drought conditions, with shifts in proportion towards more methanotrophs. In addition, the
427 relative abundance of methanogens and methanotrophs increased – with large increases observed
428 within genera with expanded genomic features that enable better tolerance towards oxidative
429 environments.

430

431 **4.1 Structural shifts in response to drought**

432

433 CH₄ emissions from peatlands have been largely attributed to the metabolism of methanogens,
434 balanced by oxidation via methanotrophs (Dean et al., 2018). Here, we hypothesised that the
435 proportion of methanogenic community shifts towards more methanotrophic relative abundances
436 when exposed to drought conditions. This hypothesis was confirmed when we observed an increase
437 of 8% in the relative abundance of methanotrophs during drought. The proportion of methanogens
438 reduced by 8% in favor of methanotrophs under drought conditions. This shift was also reflected
439 in the CH₄ flux, where both *C. vulgaris* and *R. alba* ecotypes had significantly lower fluxes during
440 the drought. However, in the *E. vaginatum* ecotype, a significant reduction was not observed,
441 despite showing the same trend as the other ecotypes. We hypothesise that the drop in water table
442 depth led to increased O₂ availability, increased phenol oxidase activity and higher peat
443 temperatures. This shift in habitable zone within the peat profile provided a new ecological niche
444 for methanotrophs and other bacteria to expand into, which were previously un-inhabitable for
445 methanotrophs due to the anoxic conditions and high concentration of phenolic compounds. This
446 result is in agreement with findings from Amodeo et al. (2018), where the authors reported that the
447 optimum growth for methanotrophs is between 20°C and 25°C combined with a 1:1 ratio of
448 CH₄:O₂. Therefore, the increased peat temperature, newly habitable oxic zone and CH₄ originating
449 from deeper anoxic layers led to increased abundances of methanotroph communities during the
450 drought.

451

452 The three main methanotrophs driving the increased proportion under the drought were type II:
453 *Methylocella*, *Methylosinus* and type I: *Methylococcus*. Type II genus *Methylocella* and
454 *Methylosinus* contributed the highest to the total microbial sum, which was also observed in
455 previous studies (Ho et al., 2011). Ho et al. (2011) observed the same pattern of rapid initial growth
456 of type II methanotrophs in severely disturbed microcosms from rice paddies, whereas growth of
457 type I methanotrophs were stunted. Although type I methanotrophs were not stunted in our results
458 (i.e., *Methylococcus*), type II methanotrophs increased more in proportion when compared to type
459 I, indicating that type II methanotrophs are highly adaptive to drought conditions and are not
460 limited by their main substrates CH₄ and O₂ and other nutrients, but rather, the amount of habitable
461 zone within the peat column.

462

463 Interestingly, the relative abundance of both methanogens and methanotrophs increase during
464 drought. Several studies concluded that methanogenesis is suppressed upon exposure to O₂ (Ma et
465 al., 2012, Yuan et al., 2009), while we observed a reduction in CH₄ flux but not a total cessation
466 of CH₄ emissions. The same was observed by Rinne et al. (2020) and similar results have been
467 observed in flooded paddies, lake sediments and bromeliad tanks where members belonging to
468 *Methanocellaceae* and *Methanosarcinaceae* increase in relative abundance following desiccation
469 (Brandt et al., 2015, Conrad et al., 2014). One explanation to the increase in relative abundance
470 can be explained by methanogens including *Methanoregula*, *Methanosarcina* and
471 *Methanosphaerula* that possess expanded genomic features that enable better adaptation to

472 oxidative environments (Lyu and Lu, 2018). In this study, the abundance of the dominant
473 methanogen, the hydrogenotrophic *Methanoregula*, remained unchanged during both drought and
474 non-drought conditions, neither decreasing nor increasing significantly. However, the most
475 metabolically versatile of the methanogens, *Methanosarcina*, increased significantly under drought
476 conditions within the ecotype *E. vaginatum*, presumably due to the presence of oxygen-detoxifying
477 enzymes such as catalase, superoxide dismutase, and superoxide reductase that give this
478 methanogen a particular eco-physiological advantage that allows for growth during desiccation
479 (Angel et al., 2011, Erkel et al., 2006, Conrad et al., 2014). Therefore, the physiological
480 characteristics of the community indicate that both hydrogenotrophic and acetoclastic
481 methanogens, especially facultative methanogen members belonging to the class
482 *Methanosarcinales*, that are more resilient to drought conditions and O₂ than other genera observed
483 here.

484
485 Methanogenesis and microbial growth are temperature dependent processes (van Hulzen et al.,
486 1999). Although methanogenesis can be inhibited when exposed to O₂, our flux measurements
487 show that not all methanogenesis stopped. It is possible that, following the drop in water table, not
488 all water evaporated, resulting in peat macropores acting as anoxic microbial refuges. It is also
489 possible that microscale anoxic sites are formed within soil aggregates, which provide an
490 ecological refuges for the survival of methanogens exposed to O₂ stress (Yuan et al., 2009). These
491 ecological refuges still hold the necessary environmental conditions for methanogenesis to occur,
492 but with an increased temperature that can increase metabolic activity.

493
494 Incubation studies by van Hulzen et al. (1999) found that alternative electron acceptor reduction
495 increases with a rise in temperature, indicating that available electron acceptors will be reduced
496 sooner – resulting in increased methanogen growth. When all alternative electron acceptors are
497 consumed, the population size of the methanogens is the limiting factor in CH₄ production.
498 Therefore, increased soil temperature coupled with the eco-physical advantage demonstrated by
499 multiple functional groups may explain the increased relative abundance of methanogens during
500 the drought, which is consistent with other studies (Brandt et al., 2015, Conrad et al., 2014).

501
502 During drought and the lowering of the water table, the relative abundance of aerobic
503 methanotrophs increased significantly. Previously uninhabitable anoxic environments within the
504 peat column became aerobic, allowing a competitive edge for genera such as *Methylocella* and
505 *Methylosinus*, which are strict aerobes. Previous studies have reported increased methanotroph
506 abundance associated with higher magnitudes of CH₄ flux (van Hulzen et al., 1999, White et al.,
507 2022), however the same results are not observed here, indicating that the reduction in CH₄ flux is
508 caused by an increase in methanotrophy. *Methylocella*, a facultative methanotroph, was the
509 dominant methanotroph in all ecotypes. This trend can be explained by its capability of growing
510 on CH₄ as well as on multicarbon substrates (Dedysh and Dunfield, 2011). This means that under
511 stressful conditions such as drought, *Methylocella* can metabolize via multiple alternative
512 metabolic pathways, yielding a competitive advantage over other obligate methanotrophs. These
513 results are further confirmed by Ho et al. (2015) and Ma et al. (2012), where they showed that the
514 recovery of type I methanotrophs needed more time between drying events. Therefore, during a
515 drought we expect type II methanotrophs to dominate. Our results indicate that methanotrophs are
516 highly resilient to droughts, but their resilience may still reach a ‘critical point’ where activity is
517 no longer recovered if droughts persist on longer time scales and increase in frequency.

518

519 4.2 Taxonomic diversity between ecotypes

520 During the drought, the overall diversity of methanogens and methanotrophs did not change. To
521 identify whether the overall diversity shifts following drought, and which peatland ecotypes holds
522 the highest resilience, we assessed α -diversity and between-ecotype β -diversity. We observed
523 small non-significant variation in the means of α -diversity between drought and control years.
524 These findings are in line with results from Kim et al. (2017), where there were no differences in
525 the diversity and composition of the microbial communities between control and a 4 week drought.
526 In contrast, Zhong et al. (2017) showed significant difference in observed species and Shannon α -
527 diversity of prokaryotic microbiota following water table draw down. However, these results were
528 based on a 46-year time interval where the original peatland was drained for livestock grazing.
529 These results indicate the need to research the effects of repetitive and long-term disturbance from
530 drought in the future. Therefore, we conclude that the resilience of methanogens and
531 methanotrophs to the effects of drought is high in the short term, but more research is needed on
532 the effects on community structure and function during sustained droughts.

533

534 4.3 Effect of ecotype on CH₄ fluxes

535 During the drought, significantly lower CH₄ fluxes were observed in *C. vulgaris* and *R. alba*
536 ecotypes, but not in *E. vaginatum*, although a similar trend was detected. We believe that the
537 presence of aerenchyma tissue within the sedge *E. vaginatum* tillers, plus the ability for the sedge
538 species to access anoxic layers through deep roots, allowed access to CH₄ produced in deeper
539 anoxic layers. This physiological trait facilitates the transport of CH₄ produced in deep anoxic peat
540 layers directly to the atmosphere, by-passing aerobic upper layers where methanotrophy can
541 oxidise CH₄. Our results are consistent with previous studies conducted in peatlands where
542 vegetative cover is directly related to the magnitude of CH₄ flux (Keane et al., 2021, Korrensalo et
543 al., 2018).

544

545 The ecotype with the largest reduction in CH₄ emissions during the 2018 drought was *R. alba*. This
546 ecotype is usually dominated by high water table depths and sphagnum mosses, indicative of
547 conditions favoring CH₄ production. Interestingly, previous studies conducted across the globe
548 have identified moss-associated methane oxidizers inhabiting *Sphagnum* (Kip et al., 2010), and
549 revealed that moss-associated methane oxidizers can exceed methanogenic activity in terrestrial
550 sites by up to two orders of magnitude (Liebner et al., 2011), but this relationship was not observed
551 in our results. One possible explanation is the presence of *R. alba*, a sedge species with aerenchyma
552 tissue similar to *E. vaginatum*. However, the rooting length of *R. alba* is substantially more shallow
553 than *E. vaginatum*, resulting in limited access to deep anoxic layers. Thus, we do not observe the same
554 by-passing of aerobic upper layers where methanotrophs can oxidize CH₄, therefore reducing net CH₄
555 emissions.

556

557 Finally, the *C. vulgaris* ecotype had the lowest CH₄ emissions in both 2017 and 2018, consistent
558 with previous studies (Keane et al., 2021). *C. vulgaris* dominated the drier portion of the bog,
559 where it is expected to see a combination of low methanogenesis and high methanotrophy, due to
560 the aerobic conditions and lack of plant mediated CH₄ transport through sedges.

561

562 **4.4 Functional genes during drought**

563 Our results indicate that the abundance of functional genes related to CH₄ metabolism significantly
564 change when exposed to drought conditions. 12% of the variance in abundances can be explained
565 by year, with the highest variation in functional gene abundance observed under non-drought
566 conditions. One possible explanation is that droughts increase the relative abundance of genera
567 with greater genetic capacities to survive drought conditions, since these individuals can take better
568 advantage of the aerobic conditions that inhibit metabolic activity in other genera.

569 The top three genes that contributed the most to the dissimilarity during the drought were *hdrA*,
570 carbon monoxide dehydrogenase large subunit (*cutL*) and hydrogen dehydrogenase. The *hdrA*,
571 combined with methyl-coenzyme M (*mcrA*), function together in the biological formation of CH₄.
572 *mcrA* catalyzes the conversion of methyl-coenzyme M and coenzyme B into CH₄ and the
573 heterodisulfide of coenzyme M (HS-CoM) and coenzyme B (HS-CoB) (Scheller et al., 2010,
574 Thauer, 2019). Subsequently, CoM and CoB must be reduced to regenerate the CoM-SH and CoB-
575 SH thiols that are used as electron donors by *mcrA*, which is then catalyzed by *hdrA* (Scheller et
576 al., 2010, Buan et al., 2011). White et al. (2022) observed a co-dependence between *mcrA* and
577 *hdrA*, indicating the close nature of the two genes for the biological formation of CH₄. Here we
578 observe significant increases in the relative abundance of *hdrA* – but not *mcrA* – potentially
579 resulting in a reduction in the conversion of *mcrA* and coenzyme B into CH₄. The lower fluxes
580 suggest this, but it is challenging to determine whether the abundances of *hdrA* come exclusively
581 from methanogens, as *hdrA* is a common gene shared between multiple microbial groups including
582 *Acetogens*, sulfur oxidizing *Archaea* and *Bacteria* (Ernst et al., 2021). Therefore, it is difficult to
583 determine whether the increase is related to methanogens or other microbial communities. In
584 addition, we observed significant increases in the relative abundance of carbon monoxide
585 dehydrogenase genes (*cutL*, *coxL*, *coxS* and *cutM*) under drought conditions. According to Ferry
586 (2010), it is not yet known if carbon monoxide is a viable energy source for methanogens in
587 peatland environments. The association between methanogenesis and the Acetyl-CoA pathway
588 appears to be much more flexible than previously thought (Borrel et al., 2016).

589

590

591 **4.5 Taxonomic diversity**

592 In 2017, the mean α -diversity was $2.30 (\pm 0.24)$, while in 2018 the mean α -diversity was 2.49
593 (± 0.25) , but this is a non-significant difference. Between ecotypes, diversity was determined with
594 the β -diversity index (Figure 7). During 2017, the highest mean distance of group members to the
595 group centroid was observed in *R. alba* plots (0.36 ± 0.26) followed by *E. vaginatum* (0.33 ± 0.11)
596 and *C. vulgaris* (0.33 ± 0.13). In 2018, this order was not observed and the *C. vulgaris* ecotype
597 had the highest mean β -diversity (0.36 ± 0.0), followed by *R. alba* (0.25 ± 0.09) and *E. vaginatum*
598 (0.17 ± 0.09). Although we see small shifts between ecotypes and years, the ANOVA's p-value is
599 not significant, meaning that group dispersions are homogenous ($p \geq 0.05$).

600

601 **5 Conclusions**

602

603 Our study provides in-situ evidence on how drought affects the functional potential
604 of microorganisms responsible for CH₄ production and oxidation in hemi-boreal peatlands. The
605 functional potential during the drought differed significantly when compared to the previous year.
606 In response to the drought, the proportion of methanogens to methanotrophs shifted in favor of
607 methanotrophs, driven by the facultative *Methylocella*. Our results suggest that (I) specific
608 functional groups respond differently to drought events due advantageous genomic traits giving a
609 competitive edge when under oxidative stress, (II) the diversity of methanogens and
610 methanotrophs does not change under drought and that not one type of ecotype holds the best
611 ecological niche, and finally, (III) peatlands dominated by sedge species *E. vaginatum* yield the
612 highest fluxes of CH₄ under drought conditions. To be able to predict the effect of anthropogenic
613 climate change, including drought events more accurately on methanogen and methanotroph
614 communities, additional attention should be paid towards the frequency and length of drought
615 events. We observe a highly resilient methanogen community, which surprisingly expanded in
616 relative abundance during drought conditions, but this increase is not reflected in CH₄ emissions
617 presumably due to the even higher increase in methanotrophy. This high abundance of both
618 communities indicates that the severe drought from 2018 did not deteriorate the functional
619 potential of the peatland ecosystem to emit CH₄ to the atmosphere.

620

621 **Acknowledgments**

622

623 We would like to thank the staff from the Centre for Genomic Research, Liverpool and Roche
624 Diagnostics Scandinavia for probe hybridization and sequencing. Special thanks go to our
625 colleagues, friends and family for the valuable discussions, and a special thank you to field
626 assistants Oskar Ström and Thomas Golin for working hard during 2018 the heatwave. This study
627 has been made possible by the Swedish Infrastructure for Ecosystem Science (SITES), at the
628 Skogaryd research catchment station. SITES receives funding through the Swedish Research
629 Council under the grant no 2019/8-365 and 2017-00635. Data handling was enabled by resources
630 in project SNIC 2019/8-365 provided by the Swedish National Infrastructure for Computing
631 (SNIC) at UPPMAX.

632

633 **Open Research**

634 The annotated seq

635 uence data used for analysing the functional potential of the microbial community in the study are
 636 available at the MG-RAST repository via project ID: 91145 with open access to public. The
 637 supplement related to the article is available online at
 638 <https://zenodo.org/record/7472945#.Y6RT9xXMJaQ>.

639

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821

822 **Figure 1:** (a) Location of Mycklemossen (black star) within Sweden (b) aerial photo of
823 Mycklemossen, black square shows the location of the sampling area and (c), map of
824 peat core and ancillary measurement locations. Map data sources: © EuroGeographics
825 and © Lantmäteriet.

826

827 **Figure 2:** A) Daily mean temperature measured in °C and B) Daily summed precipitation
828 measured in millimeters at Mycklemossen mire. 2017 is shown in blue and 2018 in red.
829 The shaded area indicates the growing season from May 1st to September 30th.

830

831 **Figure 3:** Daily mean soil temperature measured at 6 cm and 30 cm below the peat
832 surface. Blue lines indicate 2017 values while red lines indicate 2018 values. The shaded
833 area indicates the growing season from May 1st to September 30th.

834 **Figure 4:** Water table depth in meters below the surface for *C. vulgaris*, *E. vaginatum*
835 and *R. alba* ecotypes from automated sensors at the site during 2017 (A) and 2018 (B).
836 The shaded area indicates the growing season from May 1st and September 30th.
837

838 **Figure 5:** Boxplots of daily mean CH₄ flux measured during the growing season (May to
839 October) in 2017 (blue) and 2018 (red) at Mycklemossen mire. The boxes show quartiles
840 and the median, the whiskers denote data within 1.5 times of the interquartile range.
841 Colored stars denote outliers while significant differences using linear mixed effects
842 models are labelled with the p value and “n.s.” indicates a non-significant result.
843

844 **Figure 6:** Non-metric Multidimensional Scaling (NMDS) of taxonomic abundances using
845 Bray-Curtis distances. Small dots indicate individual samples while the largest dots
846 indicate the mean of all samples. Samples were analyzed at genus level and colored by
847 year 2017 (n = 17) and 2018 (n = 10).
848

849 **Figure 7:** Principal Coordinates Analysis (PCoA) of taxonomic abundances using Bray-
850 Curtis distances. Small symbols indicate individual samples while the largest symbols
851 indicate the mean of all samples. Samples were analyzed at genus level and colored by
852 ecotype and sampling year (*C. vulgaris* 2017 (n = 6) 2018 (n = 2), *E. vaginatum* 2017 (n
853 = 7) 2018 (n = 4), *R. alba* 2017 (n = 4) 2018 (n = 4)).

854 **Figure 8:** Nonmetric Multidimensional Scaling (NMDS) of functional genes using Bray-
855 Curtis distances. Small dots indicate individual samples while the largest dots indicate the
856 mean of all samples. Samples were analyzed at KO level 4 and colored by year 2017 (n
857 = 17) and 2018 (n = 10).

858 **Table 1:** Results of taxonomic contribution of each species (%) to the dissimilarity
859 between each group (SIMPER analysis) between 2017 and 2018. Taxa are ranked
860 according to their average contribution to dissimilarity between years. Average
861 abundances, percentage of cumulative contribution and Permutation p-value (Probability
862 of getting a larger or equal average contribution in random permutation of the group
863 factor) are also included. A cut-off at a cumulative dissimilarity of 70% was applied (2017
864 n = 17: 2018 n = 10).